

Freie Universität Berlin
Fachbereich Geowissenschaften

Dissertation zur Erlangung des Doktorgrades der Naturwissenschaften

**TECTONICS OF THE SOUTHERN ANDEAN
INTRA-ARC ZONE (38° – 42°S)**

von Dipl.-Geol. Matthias Rainer Rosenau

Berlin, 2004

Betreuer und Erstgutachter: Prof. Dr. O. Oncken
Zweitgutachter: Prof. Dr. M. R. Handy

Datum der Disputation: 01.07.2004

ABSTRACT

This study targets endogene and exogene processes operating at the transition of the high-relief Central Andes to the low-relief Patagonian Andes. The climatic and stationary tectonic setting of the study area including the Southern Andean intra-arc zone/Main Cordillera between latitudes 38° and 42°S provides insights into ca. 380 Ma of subduction-related magmatic and tectonic processes exhumed at different crustal levels. Analysis of multimethod and multiscale datasets aims to complete the knowledge about the tectonic evolution of the Southern Andean intra-arc zone and to improve the understanding of the dynamics of active margin systems. More specifically, remote sensing data (Landsat TM imagery and air photos) in combination with digital geologic and hydrologic maps and digital elevation models were joint to a “virtual Southern Andes” GIS (Geographic Information System)-database, analyzed geomorphometrically and interpreted morphotectonically. Structural mapping, fault kinematic analysis, petrologic investigations, and isotope-geochronology were used to form a base for conceptual tectonic models. Kinematic modeling gave for the first time reasonable estimates of Neogene amounts and rates of intra-arc deformation.

The topography of the Southern Andean intra-arc zone/Main Cordillera basically reflects the southward increasing efficiency of Neogene erosion on the western, windward side of the orogen. More specifically, a southward decreasing elevation concurrent with an increasing incision and a change from V-shaped valleys to U-shaped valleys are consistent with a change from dominantly fluvial erosion in the north to dominantly glacial erosion in the south. Accordingly, the level of exhumation of the magmatic arc basement increases southward from shallow to mid-crustal levels. Erosional unloading, heavily partitioned spatially into the south of the study area and temporarily into the glacials, accounts for at least one third of the observed differential rock uplift and exhumation since the Miocene.

The tectonic history of the Southern Andean intra-arc zone/Main Cordillera includes longlasting (tens to hundreds of millions of years) periods of extension and basin formation during the Mesozoic and Oligocene - Miocene alternating with relatively short (several million years) intervals of transpression and mountain building/exhumation during the mid-Cretaceous and Late Miocene. Amounts of cross-arc shortening are one order of magnitude smaller than in the Central Andes consistent with models of orocline formation by differential shortening. The youngest increment of intra-arc deformation in the Southern Andes is represented by the Liquiñe-Ofqui Fault Zone (LOFZ) which has been active as a brittle SC-like dextral shear zone decoupling a fore-arc sliver. A kinematic model suggests that the LOFZ has accommodated ca. 84 km (+66, -28) of dextral shear since the Pliocene. This displacement is consistent with offset of regional geological markers, vertical axis rotations, and the space provided in the fore-arc by Neogene Central Andean plateau formation. The resulting displacement rate suggests that about half of the margin-parallel component of oblique convergence between the Nazca and South American plates has been partitioned into the intra-arc zone. The remaining half of margin-parallel slip has been most probably accommodated by oblique thrusting in the accretionary wedge and to a minor amount eventually by internal deformation of the fore-arc.

There is no clear temporal relationship between plate kinematic parameters and deformation of the overriding plate. This implies that active margin deformation is controlled primarily by other factors than plate kinematics. The observation that increments of transpression follow increments of crustal extension and are late-synmagmatic with respect to emplacement of arc-granitoids suggests that subduction orogeny is triggered by magmatic weakening and crustal thinning. Initiation of the LOFZ was concurrent with collision of the Chile Ridge at the southern end of the Southern Volcanic Zone of the Andes suggesting that, under favorable mechanical conditions, extensional forces associated with subduction of an active spreading center have a primary control on fore-arc sliver formation.

ZUSAMMENFASSUNG

Diese Studie untersucht endogene und exogene Prozesse am Übergang der zentralen zu den patagonischen Anden. Der klimatische und tektonische Rahmen, in dem sich das Arbeitsgebiet in der südandinen Vulkanzone/Hauptkordillere zwischen 38° und 42° südlicher Breite befindet, ermöglicht Einblicke in ca. 380 Millionen Jahre subduktionsbezogener magmatischer und tektonischer Prozesse in verschiedenen Krustenstockwerken. Die Analyse multimethodischer und multiskaliger Datensätze zielt darauf ab, das Wissen über die tektonische Entwicklung der südandinen Vulkanzone zu vervollständigen und das geodynamische Verständnis aktiver Kontinentränder zu verbessern. Im Einzelnen wurden fernerkundliche Daten (Landsat TM und Luftbilder) in Kombination mit digitalen Geländemodellen geomorphometrisch analysiert und morphotektonisch interpretiert. Strukturenlogische Kartierung, störungskinematische Analyse, petrologische und isotopen-geochronologische Untersuchungen dienten als Basis konzeptueller tektonischer Modelle. Kinematische Modellierungen ergaben erstmals zuverlässige Abschätzungen Neogener Deformationsbeträge und -raten.

Die Topographie der südandinen Vulkanzone/Hauptkordillere reflektiert prinzipiell die nach Süden hin zunehmende Effizienz Neogener Erosion auf der westlichen, windzugewandten Flanke des Orogen. Die nach Süden hin abnehmende Höhe zusammen mit der zunehmenden Eingeschnittenheit und dem Übergang von V- zu U-Tälern sind konsistent mit einem Übergang von fluivialer zu glazialer Erosion. Dementsprechend steigt die Exhumierung des magmatischen Bogens nach Süden hin von flachkrustal zu mittelkrustal an. Entlastung durch Erosion, vor allem während der Eiszeiten und im Süden des Arbeitsgebietes, hatte zu mindestens einem Drittel Anteil an der beobachteten differentiellen Hebung und Exhumierung seit dem Miozän.

Die tektonische Geschichte der südandinen Vulkanzone/Hauptkordillere beinhaltet lang anhaltende Perioden (Zehner bis Hunderte Millionen Jahre) von Extension und Beckenbildung während des Mesozoikums und im Oligozän – Miozän abwechselnd mit relativ kurzen (wenige Millionen Jahre) Intervallen von Transpression und Gebirgsbildung/Exhumierung während der mittleren Kreide und im späten Miozän. Die Beträge Vulkanzen-orthogonaler Verkürzung sind eine Größenordnung kleiner als in den Zentralanden konsistent mit Modellen der Oroklinenbildung durch differentielle Verkürzung. Das jüngste Inkrement der Deformation im Bereich der südandinen Vulkanzone wird durch die Liquiñe-Ofqui Störungszone (LOFZ) repräsentiert, die als spröde SC-ähnliche dextrale Scherzone aktiv gewesen ist und zur Abkopplung eines *fore-arc slivers* geführt hat. Ein kinematisches Modell zeigt, dass die LOFZ ca. 84 km (+66, -28) dextrale Scherung seit dem Pliozän aufgenommen hat. Dieser Versatz ist konsistent mit dem Versatz regionalgeologischer Marker, Blockrotationen und dem Platz der durch die Neogene zentralandine Plateaubildung zur Verfügung steht. Die resultierende Versatzrate deutet an, dass ca. die Hälfte der plattenrandparallelen Komponente der schiefen Konvergenz zwischen der Nazca und der südamerikanischen Platte im Bereich der Vulkanzone aufgenommen worden ist. Der verbleibende Anteil wurde vermutlich durch schiefe Akkretion und interne Deformation des *fore-arcs* aufgenommen.

Es gibt keinen klaren zeitlichen Zusammenhang zwischen plattenkinematischen Parametern und der Deformation der Oberplatte, was den Schluss nahe legt, dass die Deformation am aktiven Kontinentalrand durch andere Faktoren kontrolliert wird. Im Einzelnen folgen transpressive Inkremente auf Inkremente krustaler Ausdünnung und sind spätmagmatisch im Bezug auf die Platznahme von Granitoiden im magmatischen Bogen, was die Folgerung zulässt, dass magmatische Schwächung und krustale Ausdünnung Auslösefaktoren der Subduktionsogenese sind. Die Entstehung der LOFZ erfolgte zeitgleich mit der Kollision des Chile Rückens am südlichen Ende der südandinen Vulkanzone, was den Schluss nahe legt, dass unter geeigneten mechanischen Bedingungen Extensionkräfte im Zusammenhang mit der Subduktion eines aktiven Spreizungszentrums eine Hauptursache der *fore-arc* sliver-Entstehung ist.

OVERVIEW

1	THE SOUTHERN ANDES	1
2	THE VIRTUAL SOUTHERN ANDES DATABASE	10
3	TOPOGRAPHY	19
4	PALEOTECTONICS	33
5	NEOTECTONICS	57
6	FAULT KINEMATICS	81
7	STRAIN MODELING	95
8	GEOCHRONOLOGY	109
9	CONCLUSIONS AND GEODYNAMIC IMPLICATIONS	129

CONTENT

ABSTRACT	III
ZUSAMMENFASSUNG	IV
OVERVIEW	V
CONTENT	VI
LIST OF FIGURES	IX
LIST OF TABLES	X
1 THE SOUTHERN ANDES	1
1.1 Climatic setting	1
1.2 Plate tectonic setting	2
1.3 The Liquiñe-Ofqui Fault Zone	4
1.4 Morphotectonic units and geological overview	5
1.4.1 Offshore fore-arc/Arauco peninsula	6
1.4.2 Outer fore-arc/Coastal Cordillera	7
1.4.3 Inner fore-arc/Longitudinal Valley	8
1.4.4 Intra-arc/Main Cordillera	8
2 THE VIRTUAL SOUTHERN ANDES DATABASE	10
2.1 Digital image data	10
2.1.1 Remote sensing data	10
2.1.2 Digital elevation models	10
2.2 Image processing and analysis	12
2.2.1 The filtering technique	12
2.2.2 Hill shading	13
2.2.3 Average topography	15
2.2.4 Base level mapping	16
2.2.5 Snowline extraction	17
3 TOPOGRAPHY	19
3.1 Swath profiles	19
3.1.1 Along-watershed profile (N-S)	20
3.1.2 Cross-cordillera profiles (E-W)	21
3.1.3 Interpretation	22
3.2 Relief distribution	22
3.2.1 Maximum, minimum, and mean elevation	22
3.2.2 Modal elevation	23
3.2.3 Interpretation	24
3.3 Hypsometry	24
3.3.1 Derivation of hypsometry	25
3.3.2 Results and interpretation	25
3.4 Roughness	26
3.4.1 Slope distribution	27
3.4.2 Surface ratio	28
3.4.3 Interpretation	28
3.5 Discussion	28
3.5.1 Spatial and temporal variation of erosion	28
3.5.2 Tectonic versus climatic control on uplift	31

4 PALEOTECTONICS	33
4.1 Regional geology	33
4.2 Paleotectonic structures in the Lonquimay area (38°S)	35
4.2.1 Inversion of the Neuquén basin	35
4.2.2 Inversion of the Cura-Mallín basin	36
4.3 The Lonquimay area as an example of shallow crustal tectonics	39
4.3.1 Mechanisms and kinematics	39
4.3.2 Amount of shortening	40
4.4 Paleotectonic structures in the Liquiñe area (40°S)	43
4.4.1 Fore-arc	43
4.4.2 Neltume block	45
4.4.3 Pirihueico block	46
4.4.4 Back-arc	48
4.5 Paleotectonic structures in the Hornopirén area (42°S)	48
4.6 The Liquiñe and Hornopirén areas as examples of mid-crustal tectonics	52
4.6.1 Mechanisms and kinematics	52
4.6.2 Amount of shortening	53
4.7 Discussion	54
4.7.1 Transpression at different crustal levels	54
4.7.2 Synorogenic lateral escape	55
4.7.3 Internal versus external trigger of subduction orogeny	56
5 NEOTECTONICS	57
5.1 Lineament analysis	57
5.1.1 Methodology	57
5.1.2 Results and interpretation	57
5.2 Drainage analysis	59
5.2.1 Methodology	59
5.2.2 Regional analysis	60
5.2.3 Subarea individualization	61
5.2.4 Drainage anomalies	63
5.3 Volcanic fissures	64
5.3.1 Methodology	64
5.3.2 Results and interpretation	65
5.4 Fault scarps	67
5.4.1 Methodology	67
5.4.2 Interpretations	68
5.5 Discussion	75
5.5.1 The LOFZ: a crustal scale SC-like fault zone system?	75
5.5.2 Reactivation of pre-existing discontinuities	76
5.5.3 State of stress and stress partitioning	77
6 FAULT KINEMATICS	81
6.1 Fault kinematic analysis	82
6.1.1 Data acquisition	82
6.1.2 Processing	83
6.1.3 Data analysis	84
6.2 Results of fault kinematic analysis	86
6.3 Discussion	90
6.3.1 Temporal versus spatial heterogeneity of deformation	90
6.3.2 Along-arc variation of the deformation field	93

7 STRAIN MODELING	95
7.1 Theoretical background	95
7.2 Kinematic model	96
7.3 Sensitivity analysis	97
7.4 Other kinematic constraints	99
7.4.1 Rotations	99
7.4.2 Translation and displacement	101
7.5 Timing of deformation	103
7.5.1 Ductile shear zones	103
7.5.2 Pull-apart basins	104
7.5.3 Evidences for Pliocene - active deformation	104
7.6 Discussion	106
7.6.1 Consistency with plate kinematic parameters	106
7.6.2 Internal deformation of the fore-arc	107
8 GEOCHRONOLOGY	109
8.1 The Rb-Sr system	109
8.2 Dating migmatization and deformation in the intra-arc	110
8.2.1 Petrography	110
8.2.2 Approach	111
8.2.3 White mica chemistry	112
8.2.4 Analytical technique	115
8.2.5 Analytical results	117
8.2.6 Discussion	121
8.3 Dating magmatism and metamorphism in the back-arc	124
8.3.1 Petrography	125
8.3.2 Analytical technique and results	125
8.3.3 Discussion	127
9 CONCLUSIONS AND GEODYNAMIC IMPLICATIONS	129
APPENDIX	133
I Coverage of remote sensing data	134
II Glaciers of the study area	136
III Volcanoes of the study area	137
IV Geomorphometric data	139
V Geothermobarometric data	140
VI Whole rock chemical data	142
REFERENCES	143
DANKSAGUNG	
CURRICULUM VITAE	

LIST OF FIGURES

Fig. 1.1:	Plate tectonic setting	3
Fig. 1.2:	Morphotectonic units	6
Fig. 2.1:	Digital elevation models	11
Fig. 2.2:	Hill-shaded DEMs	14
Fig. 2.3:	Average topography map	15
Fig. 2.4:	Base level map	16
Fig. 2.5:	Distribution of modern glaciers	18
Fig. 3.1:	Profiles along and across the Main Cordillera	20
Fig. 3.2:	Subareas used for terrain analysis	23
Fig. 3.3:	Relief distribution	24
Fig. 3.4:	Hypsometry	26
Fig. 3.5:	Slope distribution	27
Fig. 3.6:	Surface ratio	28
Fig. 3.7:	Erosion and exhumation	30
Fig. 3.8:	Effect of erosionally driven isostatic uplift	32
Fig. 4.1:	Geological overview	34
Fig. 4.2:	Map of Early Cretaceous structures in the Lonquimay area	35
Fig. 4.3:	Map of Late Miocene structures in the Lonquimay area	36
Fig. 4.4:	Late Miocene folds in the northern part of the Lonquimay area	37
Fig. 4.5:	Late Miocene folds in the southern part of the Lonquimay area	38
Fig. 4.6:	Late Miocene thrusts in the Lonquimay area	38
Fig. 4.7:	Kinematic model of Late Miocene deformation in the Lonquimay area	42
Fig. 4.8:	Geological map and profiles of the Liquiñe area	44
Fig. 4.9:	Micrograph of greenschist facies mylonites in the Liquiñe area	45
Fig. 4.10:	Geological map and profile through the Liquiñe migmatite dome	47
Fig. 4.11:	Micrographs and line drawings of greenschist facies mylonites on the eastern flank of the Liquiñe dome	47
Fig. 4.12:	Pre-Andean structures of the back-arc basement	48
Fig. 4.13:	Geological map and profile of the Hornopirén area	49
Fig. 4.14:	Late Miocene structures in the Hornopirén area	50
Fig. 4.15:	Micrograph of greenschist facies mylonites in the Hornopirén area	51
Fig. 4.16:	Syntectonic granitoids in the Hornopirén area	51
Fig. 4.17:	Micrograph of a submagmatic shear zone in granitoids of the Hornopirén area	52
Fig. 4.18:	Kinematic model of Late Miocene deformation in the Liquiñe and Hornopirén areas	54
Fig. 4.19:	Paleotectonic model	54
Fig. 5.1:	Landsat TM lineaments of the Southern Andean intra-arc zone	58
Fig. 5.2:	Results of lineament analysis	59
Fig. 5.3:	Results of regional drainage analysis	61
Fig. 5.4:	Drainage network anomalies	63
Fig. 5.5:	Concept and results of analysis of volcano morphology	65
Fig. 5.6:	Vn. Lonquimay and Vn. Tolhuaca	66
Fig. 5.7:	Vn. Callaqui	67
Fig. 5.8:	Volcanic field at Lago Caburga	67
Fig. 5.9:	Morphotectonic interpretation of the Lago El Barco area	68
Fig. 5.10:	Morphotectonic interpretation of the Lonquimay area	69
Fig. 5.11:	Mesoscale graben and normal faults in the Lonquimay area	70
Fig. 5.12:	Morphotectonic interpretation of the Biobío valley	71
Fig. 5.13:	Mesoscale normal fault in the Biobío valley	72

Fig. 5.14:	Flowerstructure in the Biobío valley	72
Fig. 5.15:	Morphotectonic interpretation of the Lago Todos Los Santos area	73
Fig. 5.16:	Facetted spurs south of Vn. Osorno	74
Fig. 5.17:	Morphotectonic interpretation of the Hornopirén area	74
Fig. 5.18:	Conceptual model of the LOFZ with SC-kinematics	75
Fig. 5.19:	Neotectonic faults and fractures of the LOFZ	76
Fig. 5.20:	Mohr circle constructions for volcanic feeder dikes	78
Fig. 6.1:	Distribution of faults used for fault kinematic analysis	83
Fig. 6.2:	Angelier and Hoeppner plots (examples)	84
Fig. 6.3:	“Beachball”-plot (example)	85
Fig. 6.4:	Fluctuation plot and Mohr diagram (examples)	86
Fig. 6.5:	Results of fault kinematic analysis represented by beachballs	88
Fig. 6.6:	Results of fault kinematic analysis represented by stereoplots	89
Fig. 6.7:	Bimodal distribution of fault kinematic data due to strain partitioning	92
Fig. 6.8:	Along-arc variation of fault kinematic data	93
Fig. 7.1:	Relation between the directions of maximum tangential shear strain and the strain ellipse	95
Fig. 7.2:	Hypothetical finite horizontal sectional strain ellipse of the study area	97
Fig. 7.3:	Sensitivity analysis	99
Fig. 7.4:	Rotation during simple shear	101
Fig. 7.5:	Offset constraints	102
Fig. 7.6:	Plate kinematic constraints	107
Fig. 8.1:	Migmatite versus mylonite	111
Fig. 8.2:	Alkali versus silica plots	114
Fig. 8.3:	Rb-Sr isotope systematics	120
Fig. 8.4:	Rb-Sr isochron plots	127
Fig. A1:	Coverage of remote sensing data	134
Fig. A2:	Example of high resolution air photo	135
Fig. A3:	Glacier Grey	136
Fig. A4:	Vn. Lonquimay	137
Fig. A5:	Normal fault in the edifice of Vn. Llaima	137

LIST OF TABLES

Tab. 6.1:	Results of fault kinematic analysis	87
Tab. 8.1:	Microprobe analyses of white mica	113
Tab. 8.2:	Rb-Sr data of migmatites and mylonites	119
Tab. 8.3:	Rb-Sr data of back-arc basement	126
Tab. A1:	Distribution of modern glaciers and snow fields	136
Tab. A2:	Volcanoes of the study area	138
Tab. A3:	Geomorphometric data	139
Tab. A4:	Al-in-hbl geothermobarometric data	140
Tab. A5:	Whole rock chemical data	142