



Benchmarking in a rotating annulus: a comparative experimental and numerical study of baroclinic wave dynamics

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Abstract

The differentially heated rotating annulus is a widely studied tabletop-size laboratory model of the general mid-latitude atmospheric circulation. The two most relevant factors of cyclogenesis, namely rotation and meridional temperature gradient are quite well captured in this simple arrangement. The radial temperature difference in the cylindrical tank and its rotation rate can be set so that the isothermal surfaces in the bulk tilt, leading to the formation of baroclinic waves. The signatures of these waves at the free water surface have been analyzed via infrared thermography in a wide range of rotation rates (keeping the radial temperature difference constant) and under different initial conditions. In parallel to the laboratory experiments, five groups of the MetStröm collaboration have conducted numerical simulations in the same parameter regime using different approaches and solvers, and applying different initial conditions and perturbations. The experimentally and numerically obtained baroclinic wave patterns have been evaluated and compared in terms of their dominant wave modes, spatio-temporal variance properties and drift rates. Thus certain “benchmarks” have been created that can later be used as test cases for atmospheric numerical model validation.

Keywords: benchmark, baroclinic instability, rotating annulus

1 Introduction

In the endeavor to improve weather forecasting and climate prediction techniques, the validation and fine-tuning of numerical models of large-scale atmospheric processes play clearly crucial roles. However, in such a complex system as the real atmosphere, validation tests are especially difficult to perform. Besides the issues that arise due to coarse-graining – a central problem of the numerical modeling of any hydrodynamic problem – in the case of atmospheric processes the unavoidable imperfection of the governing equations themselves is also a considerable source of inaccuracies. In the commonly applied hydro-thermodynamic equations the unresolved (or even physically not properly understood) processes

are either neglected or taken into account via empirical parametrization. Thus, the separation of discretization errors from the ones originating from the theoretical formulation of a given model poses a real challenge to researchers.

Yet, there is a way to carry out systematic and reproducible tests under controlled circumstances, and to capture a large segment of the complexity of these large-scale flows through relatively simple, tabletop-size experiments, based on the principle of hydrodynamic similarity. Under laboratory conditions it is possible to adjust the governing physical parameters and thus to separate different processes that cannot be studied independently in the real atmosphere. Therefore, laboratory experiments provide a remarkable test bed to validate numerical techniques and models aiming to investigate geophysical flows. This was one of the primary goals of the German Science Foundation’s (DFG) priority program MetStröm. Research focuses on the theory and methodology of multiscale meteorological fluid me-

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chanics modelling and accompanying reference experiments supported model validation.

One of these reference experiments was the differentially heated rotating annulus. This classical apparatus to study the basic dynamics of the mid-latitude atmosphere has been introduced by FULTZ et al. (1959) based on the principles first suggested by VETTIN (1857). The two most relevant factors of cyclogenesis, namely the planetary rotation and the meridional temperature gradient are quite well captured in this simple arrangement. The set-up (Fig. 1) consists of a cylindrical gap mounted on a turntable and rotating around its vertical axis of symmetry. The inner side wall of the annulus is cooled whereas the outer one is heated, thus the working fluid experiences a radial temperature gradient. At high enough rotation rates the isothermal surfaces tilt, leading to baroclinic instability. The extra potential energy stored in this unstable configuration is then converted into kinetic energy, exciting drifting wave patterns of temperature and momentum anomalies. The basic underlying physics of such baroclinic waves has been subject of extensive theoretical (EADY, 1949; LORENZ, 1963; MASON, 1975), numerical (WILLIAMS, 1971; MILLER and BUTLER, 1991; VON LARCHER et al., 2013) and experimental (FRÜH and READ, 1997; SITTE and EGBERS, 2000; VON LARCHER et al., 2005; HARLANDER et al., 2012) research throughout the past decades. Furthermore, some studies focused on the quantitative comparison of temperature statistics (GYÜRE et al., 2007) and propagation dynamics of passive tracers (JÁNOSI et al., 2010) obtained from annulus experiments and from actual atmospheric data. Even meteorological data assimilation techniques (RAVELA et al., 2010; YOUNG and READ, 2013) and techniques operational in meteorological ensemble prediction (YOUNG and READ, 2008; HARLANDER et al., 2009; HOFF et al., 2014) have also been studied by using annulus data.

The experimental part of the present study was conducted in the fluid dynamics laboratory of the Brandenburg Technical University at Cottbus-Senftenberg (BTU CS). The infrared thermographic snapshots of the drifting baroclinic waves at the free water surface have been analyzed in a wide range of rotation rates (keeping the radial temperature difference constant) and under different initial conditions. In parallel to the experiments, five numerical groups of the MetStröm collaboration (Goethe University Frankfurt, University of Heidelberg, FU Berlin, TU Dresden and TU Munich) have conducted simulations in the same parameter regime using different numerical approaches, solvers and subgrid parametrizations, and applying different initial conditions and perturbations for stability analysis. The obtained baroclinic wave patterns have been evaluated through determining and comparing their statistical variance properties, drift rates and dominant wave modes. Thus certain “benchmarks” are created that can be used as test cases for atmospheric numerical model validation in the future.

Similar comparative studies of experiments and numerical simulations in baroclinic annuli stretch back to the 1980s (JAMES et al., 1981; HIGNETT et al., 1985; READ et al., 1997; READ, 2003; RANDRIAMAMPINANINA et al., 2006; READ et al., 2008). In these works the comparisons were mostly based on pointwise sub-surface temperature time series. The very same experimental apparatus that was used in the present work has already been used to test and validate subgrid-scale parametrization methods of two of the numerical models also used here (see the paper of BORCHERT et al. (2014) in the present issue). In another recent comparative study, the effect of the addition of a sloping bottom topography to this set-up was analyzed both experimentally and numerically (VINCZE et al., 2014). However, to the best of our knowledge, the present study is the very first to systematically compare different numerical schemes and two series of experiments with different initial conditions.

Our paper is organized as follows. Section 2 outlines the experimental set-up, and the experimental and numerical methods used. The results are presented in Section 3. In Section 4 we summarize the results and discuss their implications on the physics of the underlying dynamics.

2 Methods

2.1 Experimental apparatus and procedures

The laboratory experiments of the present study have been conducted in the baroclinic wave tank of BTU CS. This tank was mounted on a turntable, and was divided by coaxial cylindrical sidewalls (Fig. 1) into three sections. The innermost compartment (made of anodized aluminum) housed coolant pipes in which cold water was circulated. The temperature in this middle cylinder was monitored via a digital thermometer and kept constant by a thermostat with a precision of 0.05 K. The outermost annular compartment contained heating wires and water as heat conductive medium. Here four thermometers (identical to that of the middle cylinder) provided temperature data for a computer-controlled feedback loop to maintain constant temperature; for the technical details on the applied control methods we refer to the paper of VON LARCHER et al. (2005). The temperatures in the inner and outer sections were set to the values of $18.5 \pm 0.25^\circ\text{C}$ and $26.5 \pm 0.25^\circ\text{C}$, respectively, yielding a radial temperature difference of $\Delta T = 8 \pm 0.5\text{ K}$.

The working fluid – de-ionized water – occupied the annular gap ranging from $a = 4.5\text{ cm}$ to $b = 12\text{ cm}$ in the radial direction. The water depth was set to $D = 13.5\text{ cm}$ in all experimental runs, thus the vertical aspect ratio of the cavity was $\Gamma = D/(b - a) = 1.8$. The water surface was *free* to enable the observation of surface temperature patterns via infrared thermography (the observed wavelength band is generally absorbed by glass or acrylic,

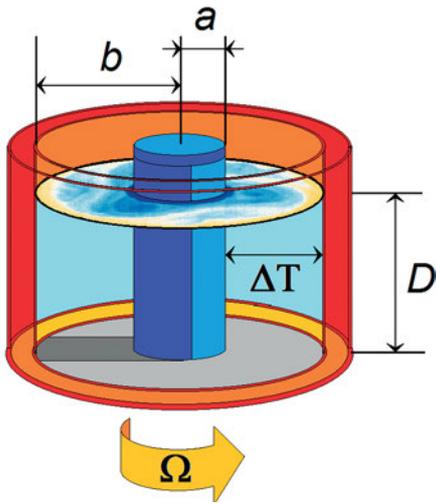


Figure 1: Schematic drawing of the laboratory set-up. For the values of the geometric parameters shown, see the text. The counter-clockwise direction of rotation is indicated.

thus covering the tank with a rigid lid was not possible). The physical properties of the fluid are characterized by its kinematic viscosity $\nu = 1.004 \times 10^{-6} \text{ m}^2/\text{s}$ and its thermal conductivity $\kappa = 0.1434 \times 10^{-6} \text{ m}^2/\text{s}$, yielding a Prandtl number of $Pr \equiv \nu/\kappa \approx 7.0$.

Since the temperature difference ΔT as well as the aforementioned geometric and material quantities were kept constant throughout the experiments, rotation rate (i.e. angular velocity) Ω was the single control parameter to be adjusted between the subsequent runs. The minimum rotation rate investigated was $\Omega_{\min} = 2.26 \text{ rpm}$ (revolutions per minute), where the flow was found to be axially symmetric, i.e. its radial and vertical structure was independent from azimuthal angle θ , indicating the absence of baroclinic instability. The highest investigated rotation rate was $\Omega_{\max} = 20.91 \text{ rpm}$. Here, four-fold symmetric baroclinic wave patterns were observed (see the exemplary thermographic snapshots of Fig. 2). Within the interval ranging from Ω_{\min} to Ω_{\max} , our measurements were taken at 17 different rotation rates. For each of these cases, two types of initial conditions were applied: the so-called “spin-up” and “spin-down” sequences. In the former (latter) initialization procedure the target rotation rate Ω was approached starting from the previously studied smaller (higher) rotation rate Ω_i , with $|\Omega - \Omega_i| \approx 1 \text{ rpm}$. The rotation rate was then gradually increased (decreased) by $\delta\Omega \approx 0.1 \text{ rpm}$ in every 2 minutes; thus, it took approximately 20 minutes to reach the required Ω from Ω_i . 10 minutes after arriving at Ω the data acquisition started and lasted for 40 to 80 minutes in each case. Afterwards this gradual increasing (decreasing) procedure of the rotation rate continued in order to reach the next Ω , with the previous parameter point as Ω_i , providing a long “spin-up” (“spin-down”) experiment series. Thus, in total 17×2 measurements were performed and evaluated. Note, that in order to enable the standard initialization procedures at the

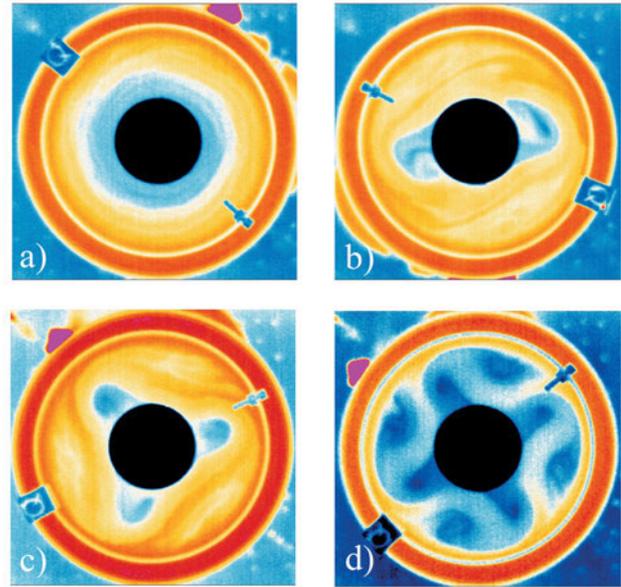


Figure 2: Four typical thermographic snapshots of surface temperature patterns in the rotating annulus. a) An axially symmetric ($m = 0$) pattern at $\Omega = 2.28 \text{ rpm}$; b) A two-fold symmetric ($m = 2$) baroclinic wave at $\Omega = 3.23 \text{ rpm}$; c) $m = 3$ at $\Omega = 4.20 \text{ rpm}$; d) $m = 4$ at $\Omega = 6.16 \text{ rpm}$ ($\Delta T = 8 \text{ K}$).

end parameter points Ω_{\min} and Ω_{\max} , the initial value Ω_i was set smaller than Ω_{\min} or larger than Ω_{\max} , when required. However, no data acquisition took place at these “out-of-range” parameter points.

The infrared camera was mounted above the middle of the tank and was fixed in the laboratory frame (not co-rotating). In every $\Delta t = 2 \text{ s}$, 640×480 -pixel thermographic snapshots were taken, providing a precision of around 0.03 K for temperature differences. The obtained temperature fields can be considered surface temperature patterns, since the penetration depth of the applied wavelength into water is measured in millimeters. The captured snapshots were acquired and stored by a computer, where they were converted to ASCII arrays (by organizing the temperature values from all pixels into matrix format) for further evaluation.

The most important classic non-dimensional parameters widely used to compare the results obtained from different baroclinic annulus set-ups are the Taylor number Ta and thermal Rossby number Ro_T (also known as Hide number). The former is basically a non-dimensional measure of rotation rate Ω and reads as

$$Ta = \frac{4\Omega^2(b-a)^5}{\nu^2 D}, \quad (2.1)$$

whereas Ro_T expresses the ratio of the characteristic velocity of the thermally driven flow to the rotation rate in the form of

$$Ro_T = \frac{Dg\alpha\Delta T}{\Omega^2(b-a)^2}, \quad (2.2)$$

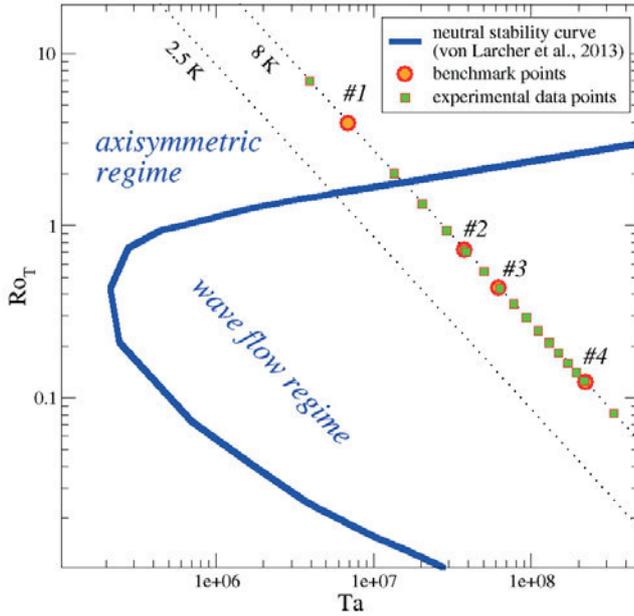


Figure 3: The neutral stability curve (thick blue line) in the parameter plane of Ta and Ro_T , as obtained via linear stability analysis by VON LARCHER et al. (2013), using the geometrical and material parameters of the BTU C-S wave tank. The dotted line corresponding to the studied radial temperature difference $\Delta T = 8\text{ K}$ is also indicated, along with the experimental data points (squares) and the benchmark parameter points (circles) of the present comparative study (see text).

where g is the acceleration due to gravity and $\alpha = 2.07 \times 10^{-4}\text{ K}^{-1}$ represents the volumetric thermal expansion coefficient of the working fluid. Note, that when the horizontal temperature difference ΔT is comparable to the vertical one, i.e. $\Delta T_z \approx \Delta T$ holds (a fairly good assumption in the present study), (2.2) corresponds to the Burger number B , defined as the squared ratio of the Rossby deformation radius $R_d = \sqrt{gd\alpha\Delta T_z}/\Omega$ to the gap width $b - a$, $B = gd\alpha\Delta T_z/(\Omega^2(b - a)^2)$.

In the case of the present study where the experiments were conducted at a practically constant value of ΔT , an inverse proportionality $Ro \propto Ta^{-1}$ holds; thus either one of these parameters per se sufficiently describes the applied conditions. Nevertheless, to demonstrate the broader context of the studied domain, we present a conceptual $Ta - Ro_T$ regime diagram in Fig. 3. The anvil-shaped thick (blue) curve represents the layout of the so-called neutral stability curve (as obtained numerically by VON LARCHER et al. (2013), to the left of which the flow is axially symmetric (radial “sideways convection”). To the right of the curve, the emergence of steady baroclinic wave patterns (as the ones in Fig. 2b, c and d) characterizes the flow, which – for even higher values of Ta – become irregular in shape as the system approaches geostrophic turbulence (a state not studied in the present paper). The curve corresponding to the constant radial temperature difference $\Delta T = 8\text{ K}$ that lay within the focus of the present work is also indicated (by

a dotted curve), along with the experimental parameter points and the four benchmark points (to be addressed later).

2.2 Numerical methods

In this subsection we briefly describe the different numerical models and methods used for the numerical simulations.

2.2.1 Governing equations and general numerical properties

The applied numerical models computed approximate solutions of the hydrodynamic equations of motion in the Boussinesq approximation (VALLIS, 2006), using different initialization procedures, grids, time steps, boundary conditions and sub-grid-scale parametrization schemes. The overall geometric parameters of the simulation domain corresponded to the aforementioned dimensions of the annular cavity of the laboratory set-up. The governing equations themselves read as:

$$\frac{\partial \vec{u}}{\partial t} + (\vec{u} \cdot \nabla) \vec{u} = -2\Omega \vec{e}_z \times \vec{u} + \frac{\delta\rho}{\rho_0} \Omega^2 r \vec{e}_r - \frac{1}{\rho_0} \nabla p + \frac{\delta\rho}{\rho_0} g \vec{e}_z + \nu \nabla^2 \vec{u}, \quad (2.3)$$

$$\nabla \cdot \vec{u} = 0, \quad (2.4)$$

$$\frac{\partial T}{\partial t} + (\vec{u} \cdot \nabla) T = \kappa \nabla^2 T, \quad (2.5)$$

where \vec{e}_r and \vec{e}_z denote the unit vectors in the radial and vertical directions (pointing upwards), respectively, \vec{u} represents the velocity field, p is the pressure and $\delta\rho$ denotes the difference between the density of the given fluid parcel and the reference density ρ_0 (in the Boussinesq approximation $|\delta\rho| \ll \rho_0$ holds). The first and second terms on the right hand side of (2.3) account for the Coriolis and centrifugal forces, respectively, which – being inertial forces – appear in the co-rotating reference frame. This form of the equation was used in the implementations of the cylFloIt, EULAG, and LESOCC2 models. In INCA the centrifugal term was omitted, since it is generally negligible in the investigated parameter range. For HiFlow³ the governing equations were solved in the non-rotating “laboratory frame”, hence there both inertial force terms were absent and the rotation of the tank entered the dynamics through the boundary conditions.

In all codes, the boundary conditions for the temperature were isothermal at the inner and outer sidewalls of the cavity (i.e. at radii $r = a$ and $r = b$). The corresponding temperatures are denoted with $T|_{r=a} \equiv T_a$ and $T|_{r=b} \equiv T_b$, respectively, yielding $\Delta T \equiv T_b - T_a = 8.0\text{ K}$, in agreement with the laboratory set-up. On the top ($z = D$) and bottom ($z = 0$) boundaries no-flux conditions were applied for the temperature (i.e. $\nabla T \cdot \vec{e}_z|_{z=0,D} \equiv 0$). For the velocities at the bottom and lateral sidewalls, the working fluid was assumed to co-rotate with the

Table 1: Summary of the basic properties of the applied numerical models.

	cylFloit	EULAG	HiFlow ³	INCA	LESOCC2
reference frame	co-rotating	co-rotating	non-rotating	co-rotating	co-rotating
centrifugal term?	yes	yes	no	no	yes
Euler acceleration?	yes	no	no	no	no
Initialization procedure	spin-up/spin-down, with dynamic $\Omega(t)$ from the end state of the previous run (at different Ω), or from non-rotating axisymmetric basic state	rotation turned on instantly	firstly, stationary eqs. solved with increased ν and κ , which then are set instantly to physically correct values	rotation turned on instantly	rotation turned on instantly, with initial state “inherited” from the final state of the previous run (different Ω)
temperature boundary conditions	$T_a = 24\text{ }^\circ\text{C}$ $T_b = 32\text{ }^\circ\text{C}$	$T_a = 16\text{ }^\circ\text{C}$ $T_b = 24\text{ }^\circ\text{C}$	$T_a = 20\text{ }^\circ\text{C}$ $T_b = 28\text{ }^\circ\text{C}$	$T_a = 16\text{ }^\circ\text{C}$ $T_b = 24\text{ }^\circ\text{C}$	$T_a = 23.5\text{ }^\circ\text{C}$ $T_b = 31.5\text{ }^\circ\text{C}$
	isothermal at the sidewalls (T_a : inner, T_b : outer), no-flux at the top and bottom boundaries ($\nabla T \cdot \vec{e}_z _{z=0,D} \equiv 0$)				
velocity boundary conditions	no-slip ($\vec{u} \equiv 0$) at the bottom and sidewalls, free slip ($\vec{u} \cdot \vec{e}_z = 0$) at the top	no-slip ($\vec{u} \equiv 0$) at the bottom and sidewalls, free slip ($\vec{u} \cdot \vec{e}_z = 0$) at the top	rigid body rotation at all boundaries ($\vec{u} = r\Omega\vec{e}_\theta$)	no-slip ($\vec{u} \equiv 0$) at the bottom and sidewalls, free slip ($\vec{u} \cdot \vec{e}_z = 0$) at the top	no-slip ($\vec{u} \equiv 0$) at the bottom and sidewalls, free slip ($\vec{u} \cdot \vec{e}_z = 0$) at the top
grid type	regular cylindrical	equidistant Cartesian	cylindrical mesh	Cartesian mesh blocks	curvilinear Cartesian mesh
number of grid points	$40 \times 60 \times 50$ ($r - \theta - z$)	$192 \times 192 \times 67$ ($x - y - z$)	$21 \times 76 \times 41$ ($r - \theta - z$)	$160 \times 160 \times 90$ ($x - y - z$)	1: $76 \times 213 \times 137$ 2: $86 \times 241 \times 153$ ($r - \theta - z$)
grid points total	120 000	2 469 888	65 436	2 304 000	1: 2 217 756 2: 3 171 078
subgrid-scale parametrization	ALDM	n.a.	n.a.	ALDM	n.a.
grid spacing (min./max.)	$\Delta r : 1.88\text{ mm}$ $r\Delta\theta : 4.71/12.57\text{ mm}$ $\Delta z : 2.7\text{ mm}$	$\Delta x; \Delta y : 1.35\text{ mm}$ $\Delta z : 2.04\text{ mm}$	$\Delta r : 2.785/5.250\text{ mm}$ $r\Delta\theta : 3.720/9.921\text{ mm}$ $\Delta z : 1.700/5.625\text{ mm}$	$\Delta x; \Delta y : 1.55\text{ mm}$ $\Delta z : 0.4/1.8\text{ mm}$	1: $\Delta r : 0.6/1.4\text{ mm}$ $r\Delta\theta : 1.3/3.5\text{ mm}$ $\Delta z : 0.6/1.1\text{ mm}$ 2: $\Delta r : 0.4/1.6\text{ mm}$ $r\Delta\theta : 1.2/3.1\text{ mm}$ $\Delta z : 0.3/1.0\text{ mm}$
integration time step δt	$\langle \delta t \rangle \approx 0.1\text{ s}$ (adaptive)	0.0025 s	0.25 s	$\langle \delta t \rangle \approx 0.05\text{ s}$ adaptive in the initial phase; afterwards: $\delta t = 0.0375\text{ s}$	1: $\langle \delta t \rangle \approx 0.033\text{ s}$ 2: $\langle \delta t \rangle \approx 0.018\text{ s}$ (adaptive)
sample rate Δt	3 s; 5 s	5 s	0.25 s	5.625 s	1 s

tank (rigid body rotation). For the codes implemented in the co-rotating frame this yields no-slip conditions ($\vec{u}|_{z=0} = \vec{u}|_{r=a} = \vec{u}|_{r=b} \equiv 0$), whereas for HiFlow³ (non-rotating frame) this condition translates to $\vec{u}|_r = r\Omega\vec{e}_\theta$ for $r = a$ and $r = b$ at all depths, and for all values of r at the bottom ($z = 0$). In all codes at the “free” water surface the slip conditions ($\nabla u \cdot \vec{e}_z|_{z=D} = \nabla v \cdot \vec{e}_z|_{z=D} \equiv 0$) and $w|_{z=D} \equiv 0$ were applied (u and v being the two horizontal and w the vertical velocity components).

The boundary conditions, reference frames, grid types and sizes and other general properties of the ap-

plied numerical codes are summarized in Table 1. In the following subsections we briefly introduce these models and discuss their most important features.

2.2.2 cylFloit

The implementation of the *cylindrical flow solver with implicit turbulence model* (cylFloit) is described in BORCHERT et al. (2014). The numerical model is based on a finite-volume discretization of the governing equations on a regular cylindrical grid. The subgrid-scale

turbulence is implicitly parameterized by the Adaptive Local Deconvolution Method (ALDM), see [HICKEL et al. \(2006\)](#). Time integration is done using the explicit low-storage third-order Runge-Kutta method of [WILLIAMSON \(1980\)](#).

The temperature dependence of the density deviation $\delta\rho(T)$, kinematic viscosity $\nu(T)$ and thermal diffusivity $\kappa(T)$ was approximated in the form of second-order polynomial fits to empirical reference data for the studied temperature range. Because of this temperature dependence, ν and κ depend implicitly on space and time, which is the reason why the viscous stress and the heat conduction have slightly different forms than the right-most terms in (2.3) and (2.5). In order to simulate the spin-up and spin-down of the annulus, the Euler acceleration $-(d\Omega/dt)r$ is added to the right-hand side of the azimuthal component of (2.3).

Three series of numerical simulations have been performed by cyFloIt: the “from scratch” series (i), where the studied state at a target rotation rate Ω was reached after initializing the system from a non-rotating axially symmetric initial state; and the “spin-up” (ii) and “spin-down” (iii) series, where a rotation rate evolution $\Omega(t)$ similar to the aforementioned laboratory sequences was imitated. The numerical parameters of these simulations are listed in the second column of Table 1.

(i) The “from scratch” simulations: In this initialization procedure, firstly an axially symmetric (thus, two dimensional; 2d) stationary solution was computed within a physical time of $t_{2d} = 10800$ s (3 hrs), with $\Omega = 0$, but with the aforementioned boundary conditions. To obtain an axially symmetric solution, the number of azimuthal grid cells was set to $N_\theta = 1$, thus reducing the problem to 2d. Then, starting from this state the full 3d simulation was initialized with a spin-up from zero angular velocity to its final value Ω_f as:

$$\Omega(t) = \begin{cases} 0, & 0 \leq t \leq t_{2d} \\ \frac{\Omega_f}{2} \left\{ 1 - \cos \left[\frac{\pi}{\tau} (t - t_{2d}) \right] \right\}, & t_{2d} < t \leq t_{2d} + \tau \\ \Omega_f, & t > t_{2d} + \tau \end{cases} \quad (2.6)$$

Here Ω_f is the final constant angular velocity used in the experiment and τ denotes the spin-up period of the rotating annulus ranging from 20 s for Ω_{\min} to 910 s for Ω_{\max} ([BORCHERT et al., 2014](#)). To trigger the formation of baroclinic waves, low amplitude random perturbations were added to the temperature field, with a maximum amplitude of $\delta T_{\text{pert}} = 0.03|T_b - T_a|$. This 3d simulation took another 10800 s, so that the waves could fully develop. A subsequent integration time of 7200 s (2 hrs) at maximum was used to record the data analysed in the present work. For further information on this initialization method, we refer to [BORCHERT et al. \(2014\)](#).

(ii) Spin-up simulations: In these cases an initial angular velocity Ω_i and a final angular velocity $\Omega_f > \Omega_i$

were chosen. The time evolution of the angular velocity $\Omega(t)$ was then computed according to the formula:

$$\Omega(t) = \begin{cases} \Omega_i + \frac{\Omega_f - \Omega_i}{2} \left\{ 1 - \cos \left(\frac{\pi t}{\tau'} \right) \right\}, & t \leq \tau' \\ \Omega_f, & t > \tau' \end{cases}, \quad (2.7)$$

where τ' means the spin-up or spin-down period. The first simulation started with $\Omega_i = 0$ rpm and $\Omega_f = 2$ rpm, the second simulation used $\Omega_i = 2$ rpm and $\Omega_f = 3$ rpm, the third $\Omega_i = 3$ rpm and $\Omega_f = 4$ rpm and so forth up to the last spin-up simulation with $\Omega_i = 19$ rpm and $\Omega_f = 20$ rpm. The spin-up period was set to $\tau' = 1200$ s (20 min) in order to imitate the typical spin-up time scale of the laboratory runs. After the spin-up period the simulation took another 1800 s (30 min). Each simulation was initialized with fields from the previous simulation.

(iii) Spin-down simulations: The parameters and the procedures of the spin-down series were the same as for the spin-up runs, the only difference being that in this case $\Omega_i > \Omega_f$ holds. The first spin-down simulation was initialized with the results from the last spin-up simulation. It therefore used $\Omega_i = 20$ rpm and $\Omega_f = 19$ rpm, the next $\Omega_i = 19$ rpm and $\Omega_f = 18$ rpm, and so forth down to $\Omega_i = 2$ rpm and $\Omega_f = 0$ rpm. After the spin-down period of $\tau' = 1200$ s the simulations here took another 1200 s only, which was long enough for the flow to equilibrate.

2.2.3 EULAG

The EULAG framework is a multipurpose multi scale solver for geophysical flows, see [PRUSA et al. \(2008\)](#) for a comprehensive review. The framework formulates the non-hydrostatic anelastic fluid equations of motion, e.g., [GRABOWSKI and SMOLARKIEWICZ \(2002\)](#), that can be solved either in Eulerian flux form or in semi-Lagrangian advective form, and it allows for a number of assumptions for particular flow characteristics, specifically the compressible/incompressible Boussinesq approximation, incompressible Euler/Navier-Stokes equations, and fully compressible Euler equations. The governing partial differential equations are evaluated with a semi-implicit non-oscillatory forward-in-time (NFT) algorithm and a finite volume discretization ([SMOLARKIEWICZ, 1991](#); [SMOLARKIEWICZ and MARGOLIN, 1997](#); [SMOLARKIEWICZ and MARGOLIN, 1998](#)). EULAG has been successfully applied to a number of geophysical problems, documented by the large number of publications in the past years, see the list of publications with respect to applications on the EULAG model website at http://www.mmm.ucar.edu/eulag/pub_appl.html, ranging from cloud microscale to synoptic and global scale in atmospheric flows, as well as it was used for modeling oceanic flows. It is worth mentioning that also solar convection ([ELLIOTT and SMOLARKIEWICZ, 2002](#)), and urban flows ([SCHRÖTTLE and DÖRNBRACK, 2013](#)), were studied, moreover, EULAG has been also applied

for simulating injuries of the human brain, treating it as a viscoelastic fluid (COTTER et al., 2002). Apart from the possibility of considering particular flow characteristics as mentioned above, EULAG also provides a framework for Direct Numerical Simulation (DNS), Large Eddy Simulation (LES), and implicit LES (ILES). We here use the DNS approach.

For the purposes of the present study the general EULAG framework has been adapted as follows. The sidewalls and the end walls of the annulus were modeled with the immersed boundary approach (cf. GOLDSTEIN et al. (1993)), where fictitious body forces in the governing equation of motion are incorporated to represent no-slip boundaries which leads to a damping of the solution in an appropriate time interval. SMOLARKIEWICZ et al. (2007) gives a detailed description of the implementation of the immersed boundary approach in the EULAG flow solver. In our study, the damping parameters were set so that the motion at the boundaries was damped to zero within a single time step. The properties of the grid, the time step and the boundary conditions are summarized in the third column of Table 1.

The governing equations (2.3) to (2.5) were solved in the Boussinesq approximation on a Cartesian grid. For the $\rho(T)$ dependence a linear decrease of density with respect to temperature was assumed with volumetric thermal expansion coefficient $\alpha = 2.07 \times 10^{-4} \text{ K}^{-1}$, as given at $T_{\text{ref}} = 20^\circ\text{C}$. The Prandtl number was set to $Pr = 7$, corresponding to the properties of de-ionized water.

2.2.4 HiFlow³

HiFlow³ is a multi-purpose C++ finite element software providing tools for efficient and accurate solution of a wide range of problems modeled by partial differential equations (PDEs), cf. HEUVELINE (2010); HEUVELINE et al. (2012). It follows a modular and generic approach for building efficient parallel numerical solvers and introduces parallelism on two levels: coarse-grained parallelism by means of distributed grids and distributed data structures, and fine-grained parallelism by means of platform-optimized linear algebra back-ends (e.g. GPU, Multicore, Cell, etc.). Further information about this open source project can be found on the project's website <http://hiflow3.org/>. For the baroclinic wave tank scenario the governing equations (2.3) to (2.5) were considered in cylindrical coordinates in a non-rotating frame, thus the inertial force terms of (2.3) were not present in this implementation. The rotation of the system was hence taken into account by setting the proper boundary conditions at the lateral and bottom sidewalls for the azimuthal velocity component, as discussed at the beginning of this section. The grid properties, boundary conditions and other numerical parameters are summarized in the fourth column of Table 1. Material parameters ν and κ were set constant (with their standard values for de-ionized water at reference temperature $T_{\text{ref}} = 20^\circ\text{C}$),

and for the thermal expansion the linear form of

$$\frac{\delta\rho}{\rho_0} = -\alpha(T - T_{\text{ref}}), \quad (2.8)$$

was used with the standard value of $\alpha = 2.07 \times 10^{-4} \text{ K}^{-1}$.

For the calculation of the initial temperature and velocity fields the stationary version of the governing equations (2.3)–(2.5) were considered, i.e:

$$\begin{aligned} (\vec{u} \cdot \nabla)\vec{u} &= -\frac{1}{\rho_0}\nabla p - \alpha(T - T_0)g\vec{e}_z + \nu_i\nabla^2\vec{u}, \\ \nabla \cdot \vec{u} &= 0, \\ (\vec{u} \cdot \nabla)T &= \kappa_i\nabla^2T, \end{aligned} \quad (2.9)$$

with the aforementioned boundary conditions. For the determination of stationary solutions, increased values of thermal diffusivity and kinematic viscosity in the forms of $\nu_i = 100 \cdot \nu$ and $\kappa_i = 100 \cdot \kappa$ were applied for reasons of numerical stability. The resulting rotation-symmetric states were used to initialize the time-dependending simulations. Since the rotation rate of the tank was kept fixed, no spin-up or spin-down procedure was applied. Instead, slightly perturbed versions of the initial state were considered to investigate the initial state's influence on the developing baroclinic wave patterns. The temperature perturbation that was applied in some of the simulations is defined in terms of the maximum perturbation M [K] and azimuthal wave number k :

$$\delta T(r, \theta, z) = M \sin\left(\frac{r-a}{b-a}\pi\right) \cos(k\theta) \sin\left(\frac{z}{D}\pi\right),$$

for $r \in [a, b]$, $\theta \in [0, 2\pi]$, and $z \in [0, D]$. In the perturbed numerical simulations presented in this paper, $M = 0.25 \text{ K}$ and $k = 1$ was chosen.

The resulting stationary velocity field and the corresponding (occasionally perturbed) temperature field were used as the initial conditions \vec{u}_0 and T_0 of the time-dependent problem and a simulation time of between about 1,000 s up to 2,500 s have been considered. The governing equations are solved on a cylindrical mesh with 65,436 points based on a finite element method. Cellwise tri-quadratic velocity and temperature functions and piecewise tri-linear pressure functions were used. This type of so-called Taylor-Hood elements are known to be stable in the sense that they fulfil the inf-sup condition (BREZZI, 1974). In Table 2, an overview of the grid properties (repeated from Table 1), accompanied with the points of the degrees of freedom (DOF) of the applied finite element method are given. On this grid, the state of the discrete solution (velocity, pressure, and temperature) is described by $N = 2,084,604$ DOF at each point in time. In time, the Crank-Nicholson scheme was applied to the governing equations, resulting in a fully coupled nonlinear equation system with all N unknowns for each time step.

Table 2: Parameter overview of the grid and the Lagrange points (DOF) of the applied finite element method by HiFlow³.

	Grid points / DOF points for pressure	DOF points for velocity and temperature
Points in $r - \theta - z$	$21 \times 76 \times 41$	$41 \times 152 \times 81$
Points total	65,436	504,792
$\Delta r_{\min/\max}$	2.785/5.250 mm	1.438/2.625 mm
$r\Delta\theta_{\min/middle/\max}$	3.720/6.821/9.921 mm	1.860/3.410/4.961 mm
$\Delta z_{\min/\max}$	1.700/5.625 mm	0.850/2.813 mm

For the solution of the nonlinear problem in each time step, Newton's method was applied. In a typical time step, 2 or 3 steps of the Newton iteration were sufficient to solve the problem adequately. The linear equation system within each Newton step is assembled and solved on a High-Performance Computer system. A GMRES solver has been applied with block-wise incomplete LU preconditioner (ILU++, MAYER (2007)), which required ca. 200 iterations in a typical calculation.

2.2.5 INCA

INCA is a multi-purpose engineering flow solver for both compressible and incompressible problems using Cartesian adaptive grids and an immersed boundary method to represent solid walls that are not aligned with grid lines. INCA has successfully been applied to a wide range of different flow problems, ranging from incompressible boundary layer flows (HICKEL et al., 2008) to supersonic flows (GRILLI et al., 2012).

In the current context the incompressible module of INCA was used with an extension to fluids with small density perturbations governed by the Boussinesq equations (2.3) to (2.5) in a co-rotating reference frame, with the exception of the centrifugal term in (2.3). The governing equations are discretized by a finite-volume fractional-step method (CHORIN, 1968) on staggered Cartesian mesh blocks. The grid was equidistant in the horizontal directions and refined towards the bottom wall in the vertical direction. The domain was split into 32 grid blocks for parallel computing. For the spatial discretization of the advective terms the Adaptive Local Deconvolution Method (ALDM) with implicit turbulence parameterization was used (HICKEL et al., 2006). For the diffusive terms and the pressure Poisson solver a non-dissipative central scheme with 2nd order accuracy was chosen. For time advancement the explicit third-order Runge-Kutta scheme of SHU (1988) was used. The time step is dynamically adapted to satisfy a Courant-Friedrichs-Lewy condition with $CFL \leq 1.0$. The Poisson equation for the pressure is solved at every Runge-Kutta sub-step, using a Krylov subspace solver with algebraic-multigrid preconditioning. The general applicability of INCA in the Boussinesq approximation with ALDM as an implicit turbulence SGS model to stably stratified turbulent flows has been demonstrated in REMMLER and HICKEL (2012) and REMMLER and HICKEL (2013).

To represent the annulus geometry within Cartesian grid blocks in INCA, two cylindrical immersed boundaries were used representing the lateral sidewalls of the flow cavity. The Conservative Immersed Interface Method of MEYER et al. (2010) was employed to impose the boundary conditions (together with other basic numerical properties), that are listed in the fifth column of Table 1. The density changes with temperature were parametrized in a linear approximation, with the same value of α as for EULAG and HiFlow³.

The simulations were initialized with a stable temperature stratification. At $t = 0$ the wall temperature and the rotation were switched on instantaneously (no spin-up or spin-down was applied). As mentioned above, during the initial phase, the integration time step was adjusted dynamically and fluctuated around $\delta t \approx 0.05$ s. In the period of constant step size, its value was $\delta t = 0.0375$ s. The total physical duration of each run ranged from 750 s to 1500 s.

2.2.6 LESOCC2

The multi-purpose solver LESOCC2 (FRÖHLICH, 2006; HINTERBERGER et al., 2007) was used to solve the governing equations (2.3) to (2.5) in Cartesian coordinates from a co-rotating reference frame. The discretization method applied is a finite volume method with a collocated variable arrangement on curvilinear coordinates. For time integration a fractional step method was employed, consisting of a Runge-Kutta scheme as predictor and a pressure-correction equation as corrector (ZHU and RODI, 1992). The momentum interpolation of RHIE and CHOW (1983) was incorporated in the discretization for pressure-velocity coupling. Parallelization was realized by domain decomposition on the basis of block-structured grids and was implemented with MPI.

Similarly to cylFloit (and to the actual experiment) "spin-up" and "spin-down" sequences were conducted. The first simulation of the "spin-up" sequence was initiated from a stably stratified axially symmetric, non-rotating state. Then the rotation was switched on immediately. The next simulation at a higher rotation rate Ω was initiated analogously, but this time the final velocity and temperature fields of the preceding simulation were used as initial conditions. This procedure was repeated until $\Omega_{\max} = 20$ rpm was reached (in 8 subsequent simulations), and then the backward ("spin-down") series started, in which the runs were initiated from the final state obtained at a higher Ω , in the same manner.

The properties of the boundary conditions, as well as the other basic numerical parameters are listed in the last column of Table 1. For discretization two different non-equidistant curvilinear, body-fitted and block structured grid meshes were used. The time steps were adapted automatically due to a combined convection-diffusion criterion, and varied in the regime: $\delta t \in (0.0177; 0.0377)$ s.

2.3 Data processing

To reduce the parameter space to investigate, from the (either experimentally or numerically) obtained temperature fields close to the free water surface a path-wise temperature profile $T(\theta)$ was extracted along a circular contour at mid-radius $r_{\text{mid}} = (a + b)/2 = 8.25$ cm for each available time instant (black circle in the exemplary experimental thermographic image in Fig. 4a). In the cases where the temperature data were stored in Cartesian grids (i.e. for EULAG, INCA and for the laboratory experiment itself), linear interpolation was applied to gain equally spaced azimuthal temperature profiles (e.g. the black curve of Fig. 4b). During post-processing the data were transformed so that the azimuthal angle θ was measured clockwise from a given co-rotating point. For the experimental and HiFlow³ data – which were given in the reference frame of the laboratory – the rotation of the tank also had to be compensated to yield the appropriate co-rotating measure of θ .

As mentioned before, the experimentally observed thermal structures were treated as the temperature patterns at the water surface ($z = D = 13.5$ cm). Also in the cases of EULAG, HiFlow³ and LESOCC2 the temperature fields of the uppermost grid level were considered. For cylFloIt and INCA, however, the temperature profiles were extracted from the somewhat lower level of $z = 10$ cm.

In order to determine the dominant azimuthal wave modes, their corresponding amplitudes and drift rates (to be discussed in the next section), the temperature profiles $T(t, \theta)$ were analyzed using discrete spatial Fourier decomposition. After subtracting the mean temperature $\langle T(\theta; t) \rangle$ (averaged over the whole azimuthal domain of the contour at each time instant t), the remaining fluctuations could be expressed as amplitudes $A_m(t)$ and phases $\phi_m(t)$ of trigonometric functions with integer wave numbers $m = 1, 2, \dots$, as:

$$T(\theta; t) - \langle T(\theta; t) \rangle \approx \sum_m A_m(t) \sin(m\theta + \phi_m(t)). \quad (2.10)$$

Fig. 4b demonstrates this step, showing three (exemplarily selected) components: $m = 3$ (red), $m = 4$ (blue) and $m = 6$ (green) at a given time instant. The time series of $A_m(t)$ and $\phi_m(t)$ of the different numerical models and the experiments could then be easily compared using various standard methods of signal processing, to be discussed in the following section.

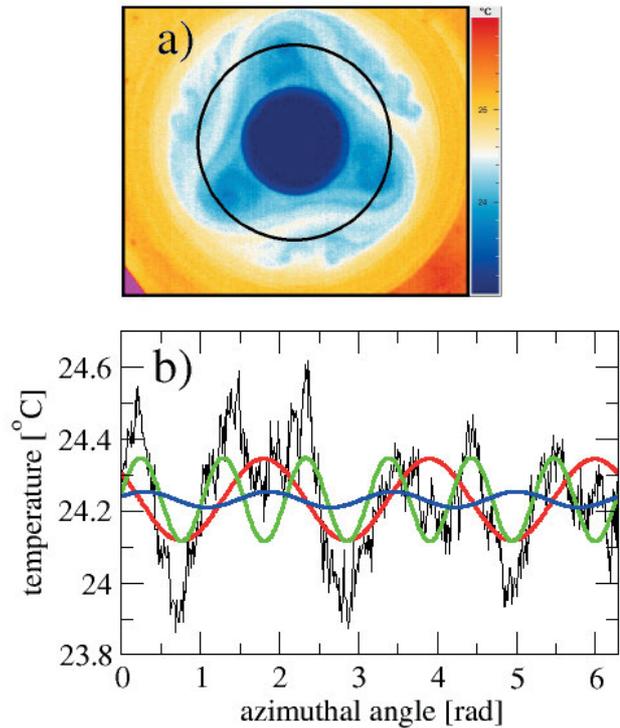


Figure 4: Three steps of data processing, demonstrated on a single thermographic snapshot of the laboratory experiment. The temperature values of the raw image are (a) extracted along a circular contour at mid-radius r_{mid} , thus the azimuthal temperature profile (b) is obtained. The Fourier components of integer wave numbers are then determined for each time step. In this exemplary case modes $m = 3, 4$ and 6 are shown by red, blue and green curves, respectively.

3 Results

3.1 Wave numbers

Firstly, the time averaged amplitudes $\langle A_m \rangle$ of the spatial Fourier components were determined in each (either experimental or numerical) case using the above described methodology. For this averaging the transient part of the wave evolution was omitted, only the quasi-stationary part of each time series was retained. To test quasi-stationarity, we sliced the time series $A_m(t)$ in question into 10–20 disjoint sections of equal length, and calculated their means and standard deviations. If the obtained statistical quantities were found to agree with each other and with those of the original long time series within a 5% margin, than the record was accepted as quasi-stationary.

The time averaged spatial Fourier spectra $\langle A_m \rangle(m)$ showed, that besides the wave number corresponding to the main azimuthal symmetry properties of a given baroclinic wave, the smaller-scale structures of the surface temperature field also leave a pronounced spectral “fingerprint”. In the Fourier space, these patterns are represented as harmonics of the basic wave number. It is to be emphasized, that the term ‘harmonic’ here is meant strictly in the sense of integer multiples of the wave number, without any further implications on the dynamics.

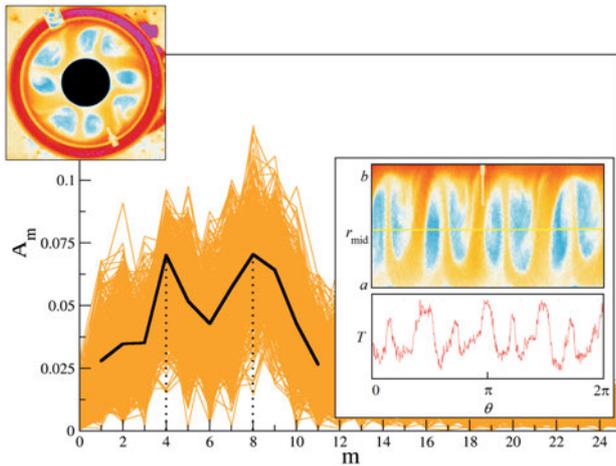


Figure 5: Spatial Fourier spectra (orange), extracted from the quasi-stationary part of a laboratory experiment ($\Omega = 17.1$ rpm, spin-down series), and their temporal average in the lower m -domain (black). In the inset, a typical thermographic snapshot is shown in polar coordinates, and the corresponding one dimensional temperature profile at r_{mid} (red curve).

3.1.1 A conceptual demonstration

As a demonstration of the physical origin of such spectral peaks, an exemplary case is shown in Fig. 5. The top left inset shows one of the original images of a given laboratory experiment, where the four-fold symmetric shape of the temperature field is apparent. The bottom right inset depicts the same image as transformed to polar coordinates: the yellow line marks mid-radius r_{mid} , and the corresponding pathwise temperature profile is also given underneath. The spatial Fourier spectra of such profiles, taken at different time instants during the same experimental run are plotted as orange curves in the main panel. Their average is also indicated (thick black curve). Manifestly, alongside the peak of $m = 4$, another significant spectral peak appears at $m = 8$, caused by the warm jet that is meandering between cold eddies (cf. insets).

In several cases among the laboratory experiments, such geometric “harmonics” even surpassed the “basic mode” in amplitude. Therefore, in order to be consistent with the traditional visual classification of wave numbers, not necessarily the largest peak was labeled the so-called *dominant wave number*. Instead, the following algorithm was applied: (i) all the significant peaks of the time-averaged spectra were determined. (ii) If two or more peaks appeared at wave numbers that are integer multiples of the first one, then the wave number m of the first peak was considered to be the dominant wave number. Even if its average amplitude $\langle A_m \rangle$ is not the largest of all, this definition still implies that the patterns bear an overall symmetry to azimuthal rotation by $2\pi/m$ (i.e. the autocorrelation of the temperature profile exhibits its largest positive peak at $2\pi/m$).

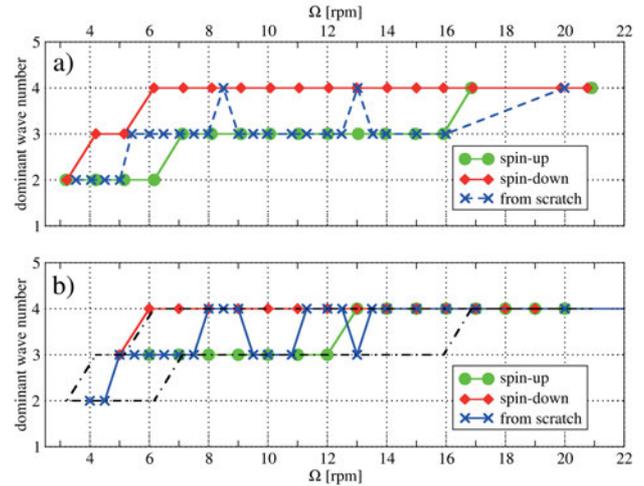


Figure 6: “Subway map” of the baroclinic annulus: the dominant wave numbers as a function of rotation rate Ω as found in the experiments (a) and in the cylFloIt simulations (b). The experimental hysteresis regime of (a) is repeated in panel (b) with dash-dotted lines.

3.1.2 The dominant wave numbers

The above defined dominant wave numbers are presented in Fig. 6a as a function of rotation rate Ω , as found in the laboratory experiments. Apparently, large hysteresis can be observed, in qualitative agreement with the findings of several previous studies (MILLER and BUTLER, 1991; SITTE and EGBERS, 2000; VON LARCHER et al., 2005), implying multiple equilibria. A broad rotation rate regime (ranging from $3.9 \text{ rpm} < \Omega < 17.1 \text{ rpm}$) exhibited different wave numbers in the “spin-up” and “spin-down” series, with $m = 3$ and $m = 4$ being the dominant modes, respectively (see green and red curves in Fig. 6). It is important to note that even in the hysteretic regime the wave patterns appeared to be stable against surface perturbations: during the experimentation process, after recording a particular pattern, irregular manual stirring was applied in the uppermost fluid layer (with penetration depth of roughly 1 cm), and afterwards, in all observed cases, the same wave pattern recovered within ca. 10 minutes of time. Despite the hysteresis, it is to be remarked, that the critical rotation rate $\Omega_{\text{crit}} \approx 3 \text{ rpm}$ of the onset of baroclinic instability and the first – critical – wave number ($m_{\text{crit}} = 2$) appeared to be unaffected by the initial conditions. We also note that all the baroclinic waves observed (except for a single transient case, encountered in one of the EULAG simulations, to be discussed later) were of the *steady wave* type, i.e. the large-scale structure of the propagating patterns did not change considerably throughout the the quasi-stationary parts of the (either experimental or numerical) runs.

To model the hysteretic behavior numerically, the cylFloIt and LESOCC2 runs imitated the experimental process via initiating the simulation of a given parameter

Table 3: Dominant wave numbers of the “benchmark” data points, as obtained in the experiment and by the numerical models. Arrows \uparrow and \downarrow mark spin-up and spin-down initial conditions, if applicable. u marks the unperturbed and p denotes the perturbed initiation states in the HiFlow³ simulations. Note, that $\Delta T = 8$ K was set constant for all the measurements, therefore the rotation rate Ω was the only variable “environmental” parameter.

notation	Ω [rpm]	experiment	cylFloit	EULAG	HiFlow ³	INCA	LESOCC2
#1	3 ± 0.2	0 – 2($\uparrow\downarrow$)	0($\uparrow\downarrow$)	2 – 3I	0(u, p)	2	0($\uparrow\downarrow$)
#2	7 ± 0.1	3(\uparrow); 4(\downarrow)	3(\uparrow); 4(\downarrow)	3	3(u); 2(p)	4	2(\uparrow); 3(\downarrow)
#3	9 ± 0.1	3(\uparrow); 4(\downarrow)	3(\uparrow); 4(\downarrow)	4	2(u); 3(p)	4	3(\uparrow); 4(\downarrow)
#4	17 ± 0.1	4($\uparrow\downarrow$)	4($\uparrow\downarrow$)	4	4(u, p)	4	3(\uparrow); 4(\downarrow)

point from the final flow state of the preceding simulation. By sequentially increasing (decreasing) the rotation rate in this manner, “spin-up” (“spin-down”) series were generated, as discussed in the previous section. Besides, the stability of the obtained state with respect to perturbed initial conditions (the analogue of manual surface stirring in the laboratory) was analyzed in the HiFlow³ simulations (the methods of perturbation are discussed in subsection 2.2.4).

The dominant wave numbers of the cylFloit runs are shown in Fig. 6b. The green and red curves represent the spin-up and spin-down series, respectively. Compared to the experimental data of panel a), the cylFloit spin-up series exhibited switches from $m = 2$ to $m = 3$ and from $m = 3$ to $m = 4$ at lower rotation rates. Nevertheless, it can be stated, that throughout the whole series, the simulations always converged to one of the experimentally observed equilibria, i.e. the cylFloit spin-up curve is enveloped by the experimental hysteresis regime (repeated in Fig. 6b with dash-dotted lines). The spin-down series, on the other hand, precisely reproduced the laboratory results, including the appropriate estimation of the critical rotation rate Ω_{crit} and critical unstable mode m_{crit} .

The dominant wave numbers obtained in an earlier experimental series (that was conducted in 2011 and had been used for the validation of the cylFloit and INCA models, see the paper of BORCHERT et al. (2014) in the present issue) are also shown in Fig. 6a in the form of a blue curve. Each of these laboratory runs had been initiated with *zero angular velocity* until the axially symmetric basic state of “sideways convection” developed. Afterwards, the rotation of the tank was accelerated so that the final rotation rate was reached within a spin-up period of ca. 20 s. In these experiments the wave patterns were observed – and remained stable – for extremely long times ranging from 6 to 12 hours after the onset of rotation. This laboratory procedure was also simulated with cylFloit (using the ‘from scratch’ strategy described in the ‘Numerical methods’ section), and the resulting data points are shown as the blue curve of Fig. 6b. It can be stated that both in the experiments and in the simulations, even though the system was initiated “from scratch” before each run, the flow occasionally converged to the states of the upper (spin-down) branch. This observation underlines the conclusion that the hysteretic regime indeed involves two distinct equilibrium

states and does not arise merely due to some slow transient phenomenon.

The experimental and numerical results for the four benchmark parameter points (for which the flow states were computed by all the numerical models) are summarized in Table 3. These points were selected to represent the three dynamical regimes observed in the laboratory: the transition zone from axisymmetric ($m = 0$) to wave flow state (#1), the hysteretic regime (#2 and #3), and the regime of higher rotation rates, where – at least in terms of the dominant wave numbers – the two branches have recombined (#4). The arrows (\uparrow and \downarrow) mark the spin-up and spin-down series, if applicable. In the case of the LESOCC2 runs, the flow states were also computed at intermediate data points (at rotation rates $\Omega = 5.5; 8.0; 10.5; 20.0$ rpm), to enable the same sequential simulation process as described for cylFloit. The data from these points, however, were not evaluated in the present study.

In the case of the HiFlow³ simulations, letters “ u ” and “ p ” denote the unperturbed and perturbed states obtained for the given rotation rate, respectively. In the “ p ”-runs additional azimuthal random perturbation was added to the initial condition (described in the previous section). In the cases of #2 and #3, the perturbed initial state led to a solution different from the unperturbed case, but no such behavior was found for #1 and #4. This is in qualitative agreement with the laboratory results, since all of these metastable states were found within the hysteretic regime (metastable, since small temperature perturbations confined to the surface region were able to change the dominant wave number in these numerical runs). It is to be noted, that the LESOCC2 and HiFlow³ models exhibited $m = 2$ at #2 in their spin-up and perturbed series, respectively, besides the (experimentally verified) $m = 3$ mode.

EULAG and INCA always converged to one of the experimentally detected states within the regime of baroclinic instability (#2 to #4). For the data point #1 close to the critical transition point, INCA found $m = 2$, and EULAG showed a dispersive transient pattern with fluctuating amplitudes at $m = 2$ and $m = 3$ (denoted with 2 – 3I in Table 3), see also VON LARCHER and DÖRNBRACK (2014) in the present issue. These may be the same type of “weak waves” that have already been observed in the BTU CS baroclinic annulus around the

critical parameter region of the onset of baroclinic instability (SEELIG et al., 2012; VINCZE et al., 2014). These findings are seemingly in contradiction to the axially symmetric solutions of the rest of the models. It is important to remark, however, that the exact experimental value of Ω_{crit} is hard to determine. At $\Omega = 2.26$ rpm the flow in the laboratory tank was clearly axially symmetric, and at the next measured data point ($\Omega = 3.19$ rpm) the first baroclinic wave pattern with $m_{\text{crit}} = 2$ has already emerged. Moreover, in the aforementioned 2011 experimental series, axially symmetric ($m = 0$) state was reported at $\Omega = 2.99$ rpm. Therefore the transition from $m = 0$ to $m_{\text{crit}} = 2$ appears to take place at $3 \text{ rpm} < \Omega_{\text{crit}} < 3.19 \text{ rpm}$, a rather narrow range.

As a general remark, it is worth to note that when a simulation does not uncover the same equilibrium as the experiment, it does not necessarily indicate a shortcoming of the simulation. Even when a variety of initialization procedures are used, the existence of multiple equilibria may not be uncovered in all cases. Also, different initialization procedures may uncover different solutions. Thus, the spin-up and spin-down simulations are more useful in this regard, i.e. they are a more robust way of uncovering the multiple equilibria.

3.1.3 Spatial harmonics and small-scale structure

Besides the dominant wave numbers, the aforementioned “harmonics” are also of relevance, as they provide a certain spectral fingerprint of the studied patterns. The wave numbers corresponding to all significant peaks of the time-averaged spatial spectra are shown in Figs. 7a and b for the laboratory experiments and for the cylFloit runs, respectively. In both panels, red crosses mark the spin-up and black circles mark the spin-down series. In each case, a peak was considered significant if its time-averaged spatial Fourier amplitude $\langle A_m \rangle$ was larger than $\bar{A} + 3\sigma_A$, where \bar{A} and σ_A are the mean amplitude and standard deviation of the whole time-averaged spectrum, respectively.

Comparing the two panels of Fig. 7, it is visible that in the laboratory experiments the presence of the harmonics was more pronounced than in the simulations. For example, the dominant wave mode $m = 4$ was always accompanied by a significant $m = 8$ in the laboratory (cf. Fig. 5), whereas it exhibited insignificant amplitudes in some of the cylFloit runs. Also, the harmonic $m = 9$ regularly appeared alongside mode $m = 3$ in the experimental data, whereas in the cylFloit results it showed up in one single case only. This mismatch might indicate that the formation of some of the eddies in the annulus (that yield the presence of these harmonics) can be caused by surface effects (e.g. wind stress, nonzero heat flux, etc.) that are not included in the numerical models.

3.2 Average temperature variance

As a measure of the overall spatial thermal variability in the azimuthal direction, the (spatial) standard deviation

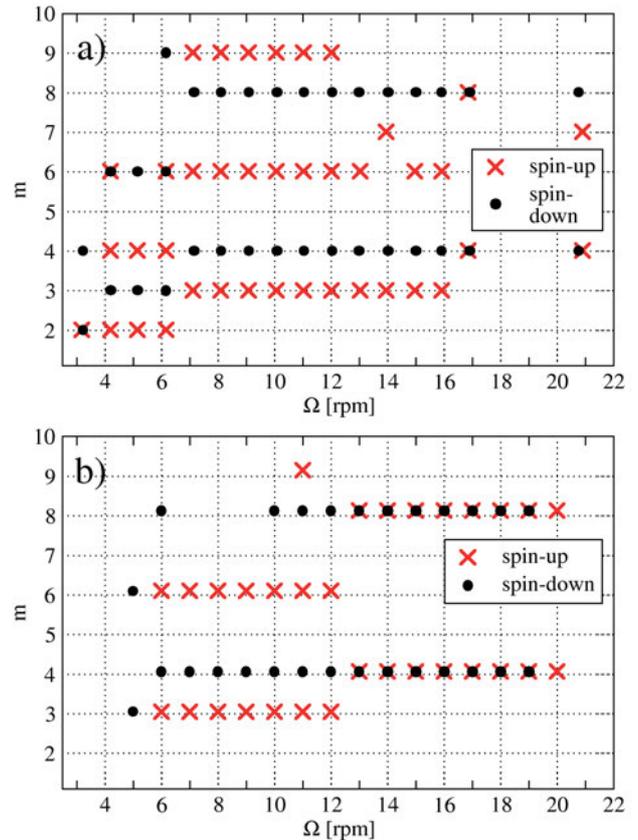


Figure 7: The distribution of significant harmonic modes in the wave number space, as a function of rotation rate Ω , as found in the laboratory experiments (a) and in the cylFloit simulations (b).

of the mid-radius temperature profile was determined at each time instant. Next, the (temporal) average of these values – denoted by $\bar{\sigma}$ – was calculated for the whole quasi-stationary part of the given (either experimental or numerical) run. The obtained values are shown in Fig. 8 as a function of the rotation rate Ω . In the graphs corresponding to those numerical simulations, where the onset of baroclinic instability was captured, this “phase transition” manifests itself in the form of a marked jump in $\bar{\sigma}$. Note, that EULAG and INCA found dominant modes of non-zero m already at the benchmark point #1, therefore in their graphs no such jump is present. Since the basic state is axially symmetric and the analyzed data were extracted from a circular contour of a constant radius r_{mid} , it is trivial that the numerical models give practically zero variance in this regime. However, due to random temperature fluctuations, the laboratory experiments (green and red curves for the spin-up and spin-down series, respectively) showed considerably larger (yet, minimum) values of $\bar{\sigma}$ in this regime.

The qualitative behavior of the spin-down experimental series in terms of $\bar{\sigma}$ is well captured by the corresponding cylFloit runs (blue curve). In both curves pronounced local maxima can be observed at $\Omega = 5$ rpm, followed by local minima at $\Omega = 6$ rpm. Both in the experiments and the cylFloit runs, this parameter point

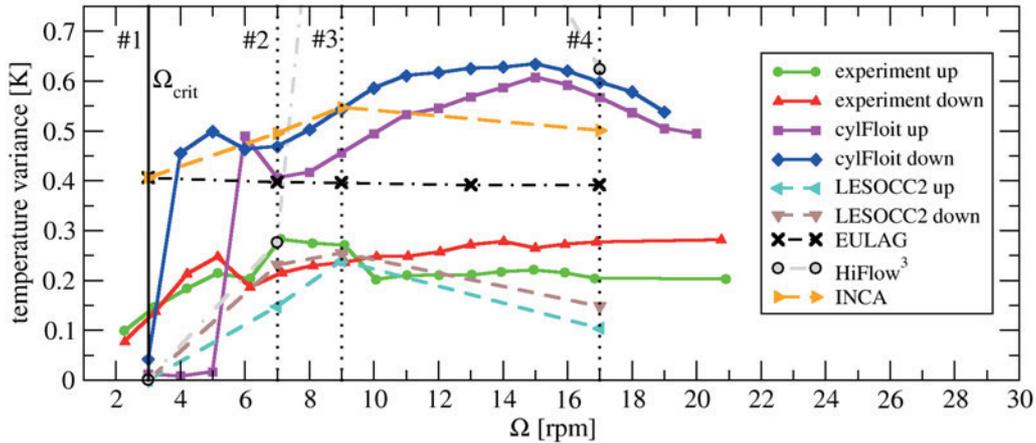


Figure 8: The average thermal variability $\bar{\sigma}$ as the function of rotation rate Ω . The numbers denote the ‘benchmark’ parameter points of Table 3.

coincides with the transition from dominant wave number $m = 4$ to $m = 3$ (as we now discuss the spin-down sequence). This may imply that the $m = 3$ patterns generally have larger amplitudes in the mid-radius section than their $m = 4$ counterparts. Thus, the reorganization of the surface pattern overrides the general decreasing trend of $\bar{\sigma}$ towards smaller values of Ω . A similar jump-wise increment is present in the experimental spin-up series as well (green curve). In this case, the transition happened at $\Omega = 7$ rpm, which, again, coincides with the transition to $m = 3$, this time from the preceding $m = 2$ state (cf. Fig. 6a). In this sequence also a similarly sharp drop of $\bar{\sigma}$ can be observed at $\Omega = 10$ rpm, which is *not* accompanied with the change of the dominant wave number $m = 3$. However, as it will be demonstrated in the next subsection, this decrease coincides with a similarly sharp change in the drift rates of the baroclinic waves, thus implying a certain state transition, even though not in terms of m .

Despite the qualitative similarity, the cylFloit and INCA runs (blue, magenta and orange curves) systematically overestimate $\bar{\sigma}$ by around a factor of 2. This, however, may well be the consequence of the fact that the temperature fields of these models are extracted from the height level of $z = 10$ cm (whereas $D = 13.5$ cm). The plotted data from LESOCC2 and EULAG (brown and black data points) on the other hand were extracted from the uppermost (surface) grid level. In terms of $\bar{\sigma}$, the former is in fairly good agreement with the experimental data, whereas the latter stays practically constant (exhibiting a minor decreasing trend only), and significantly overestimates the variance.

3.3 Drift rates

Next, the drift rates of the dominant wave modes were determined and analyzed. The discrete Fourier transform, described in the ‘‘Methods’’ section, yielded the phase shifts ϕ_m for each time instant. Thus, the quantity $\phi_m(t)/m$ could be used as a measure of the ‘‘azimuthal

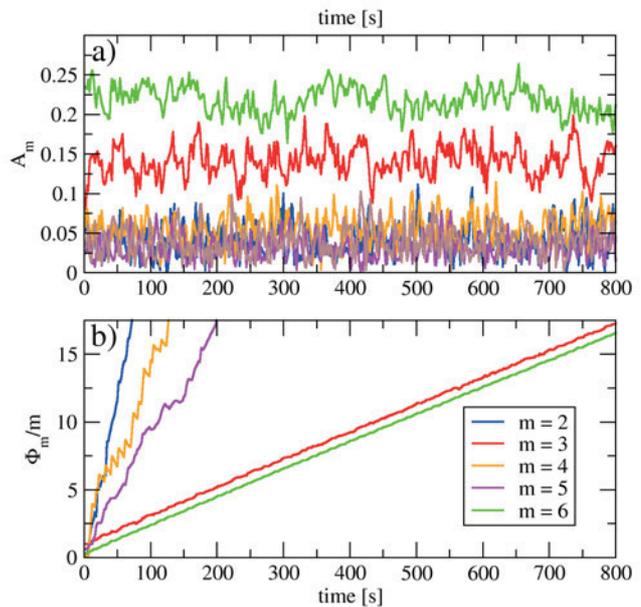


Figure 9: Temporal development of the Fourier amplitudes (a) and ‘‘azimuthal distances’’ (b) of wave modes $m = 2, \dots, 6$ in a laboratory experiment ($\Omega = 4.2$ rpm, spin-down series). Note, that the modes of the dominant wave number $m = 3$ and its ‘‘slave pattern’’ $m = 6$ – that has the largest amplitude – exhibit regular, uniform drift, whereas the small-amplitude modes provide bogus ‘non-physical’ signals in the bottom panel.

distance’’ travelled by the given component with wave number m since $t = 0$. Such time series are shown for the two largest Fourier components ($m = 3$ and $m = 6$) in the explanatory figure Fig. 9b obtained in a laboratory experiment ($\Omega = 4.2$ rpm, spin-down series), alongside with amplitudes $A_m(t)$ of the first six Fourier components in Fig. 9a. For better visualization of the evolution of $\phi_m(t)/m$ in the bottom panel, we extended the periodical $[0; 2\pi]$ range to $[0; +\infty)$ (so that the positive increments correspond to counter-clockwise propagation).

The drift rate $c_m(t)$ (angular velocity) of a given mode m could thus be obtained as the slope of the corresponding graph at time instant t , since:

$$\frac{1}{m} \frac{\partial \phi_m(t)}{\partial t} \equiv c_m(t), \quad (3.1)$$

therefore linear fits to the quasi-stationary part of the propagation could be used to determine $c_m(t)$.

It is important to mention, that, within a given experiment all the Fourier components of significant amplitudes propagated at the same drift rate, i.e. no wave dispersion was present. Consequently, although the flow pattern drifted around the annulus, its form remained unchanged. We note, that in an earlier experimental series carried out in the same set-up with the addition of *sloping bottom topography*, marked wave dispersion was observed. In that case, the stable baroclinic wave patterns emerged in the form of so-called resonant triads (VINCZE et al., 2014). Moreover, PFEFFER and FOWLS (1968) also found dispersion in their flat bottom experiment, and HARLANDER et al. (2011) reported dispersion in the wave transition region of the $Ta - Ro_T$ diagram at lower ΔT .

We compared the drift rates of the wave mode of the largest average amplitude $\langle A_m(t) \rangle$ for each run. The drift rates obtained for the laboratory experiments are presented in Fig. 10a, both for the spin-up (green) and spin-down (red) series, as a function of rotation rate Ω . An overall decreasing trend can be observed in agreement with the expectations based on quasi-geostrophic theory. Due to thermal wind balance, the velocity of the zonal background flow is expected to scale as:

$$U \propto \frac{\alpha g D \Delta T}{2\Omega(b-a)}. \quad (3.2)$$

In the linear theory of EADY (1949) the baroclinic waves themselves also propagate at the velocity of the mean flow, thus a $c_m \propto \Omega^{-1}$ scaling is to be expected. Accordingly, FEIN (1973) found in baroclinic annulus experiments the general power-law form $c_m = B(\alpha \Delta T / \Omega)^\zeta$. In the case of our experiments (the spin-down series was evaluated), these parameters were found to be $B = 4.4 \pm 0.15$ and $\zeta = 1.17 \pm 0.04$. The fit is shown in Fig. 10c – the repetition of panel a) with logarithmic scales – as a dashed line, and a $\zeta = 1$ slope proportional to the thermal wind speed is also plotted (thick grey curve). It is to be noted, that for a free-surface annulus Fein obtained $\zeta = 0.88 \pm 0.07$ (the values of B are not suitable for direct comparison between different set-ups as they depend on the actual geometrical parameters of the tanks used).

Fein also demonstrated that both in terms of factor B and exponent ζ the experiments with free surface and rigid lid exhibit significantly different scaling properties, leading to an order-of-magnitude difference between their respective drift rates (the waves in the free surface set-up being the faster). This observation underlines the extreme sensitivity of the studied system to the upper

boundary condition, and thus gives a broader context to our comparisons with the numerical results, which now follows.

The c_m values, obtained from the cylFloIt data are shown with magenta and blue curves in Fig. 10a and c, representing the drift rates in the spin-up and spin-down series, respectively. Also, the results of the “from scratch” series (always initiated from the stable $m = 0$ state) are plotted with a blue graph. Figure 10b and d show the drift rates found in the LESOCC2 (spin-up and spin-down), INCA and EULAG simulations. The general decreasing trend of drift rates was captured by the investigated models, and the drift rates of cylFloIt, INCA and LESOCC2 are in fairly good agreement with each other, yet, neither the experimentally obtained, nor the thermal wind-type scaling was reproduced by them. The drift rates are generally overestimated compared to the laboratory findings (the experimental curves and the cylFloIt “from scratch” points are repeated in 10b and d in the form of a dotted curves, and a $\zeta = 1$ power-law is also given in panel d). The EULAG simulations however – aside of the $\Omega = 3$ rpm case, where the wave pattern appeared rather irregular – were in good agreement with the experiments in terms of drift rates. The possible reasons for these differences will be discussed in the “Summary and conclusions” section.

Besides the general decreasing trend of c_m , the most marked feature in the experimental spin-up sequence (green curves in 10a and c) is a sharp drop around $\Omega = 10$ rpm, a data point which lies well within the regime of dominant wave number $m = 3$ (cf. Fig. 6a). This transition was also observed in terms of the average thermal variance $\bar{\sigma}$, as mentioned in the previous subsection.

It is to be noted, that in the experimental procedure, the discussed drop coincided with an interruption of the measurement sequence. The spin-up measurements were conducted in four campaigns on subsequent days. The measurement protocol in such cases was the following: on a new measurement day, the spin-up process was repeated from an initial axially symmetric state with a fully established sideways convection ($\Omega \approx 2$ rpm), up to the preceding data point (in this case to $\Omega = 9.1$ rpm), which was then left undisturbed for a long relaxation time (here 4 hours and 40 minutes). Afterwards, the standard spin-up procedure – described in the “Methods” section – was conducted to approach the new parameter point (in this case: $\Omega = 10.1$ rpm). Interestingly, this was the single case where the re-initiation of the measurement sequence coincided with such an abrupt change. Similar interruptions and re-initiations occurred between the data points of $\Omega = 4$ rpm and $\Omega = 5$ rpm and between $\Omega = 15$ rpm and $\Omega = 16$ rpm (and also, between $\Omega = 15$ rpm and $\Omega = 14$ rpm in the spin-down series), without any significant effect on the drift rates.

As mentioned above, the observed phenomenon was not accompanied with the change of the dominant wave number, yet, a certain *topological transition* of the surface temperature field was detected. Fig. 11 shows two

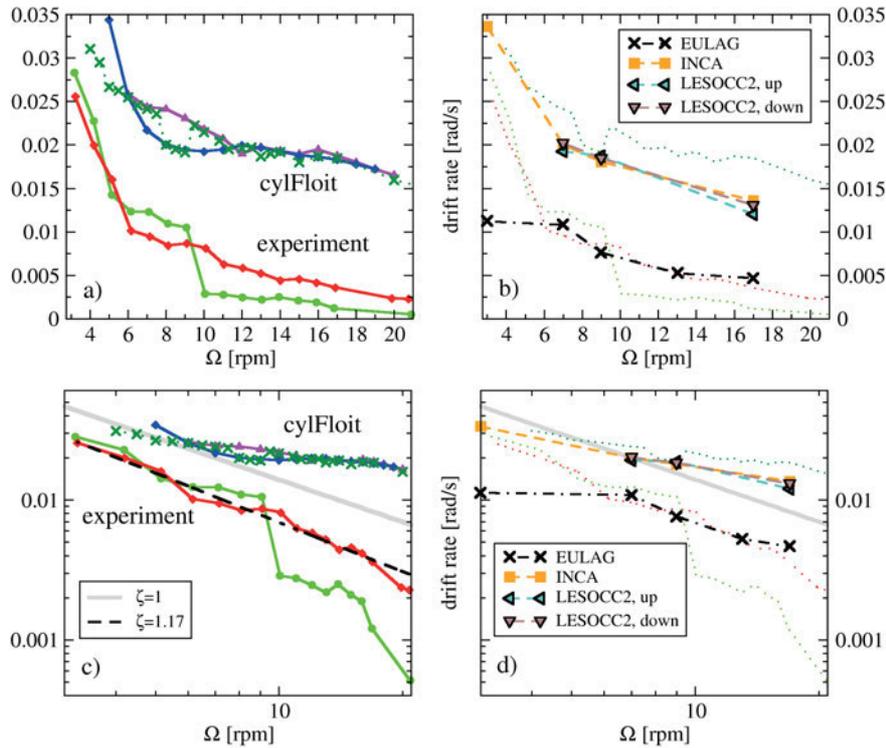


Figure 10: Drift rates of the dominant wave modes as functions of rotation rate Ω . In panel (a), the experimental spin-up (green), spin-down (red) sequences are presented, alongside the spin-up (magenta), spin-down (blue) and “from scratch” (dark green) series. In panel (b) the drift rates from other numerical models are shown. For a better comparison, three curves of panel (a) are repeated here with dotted lines, using their original color coding. The data from panels a) and b) are repeated with double logarithmic scales in panels c) and d). The power-law fit of the (spin-down) experimental data points (dashed line) and $\zeta = 1$ curves (grey) obtained via thermal wind balance are also shown.

typical snapshots, transformed to polar coordinates. The pattern characteristic to the first, “classic” type of $m = 3$ waves (observed in the range of $7.1 \text{ rpm} \leq \Omega \leq 9.1 \text{ rpm}$) is presented in panel a), whereas the structure of the slowly propagating type ($10.1 \text{ rpm} \leq \Omega \leq 15.9 \text{ rpm}$) is visible in panel b). One can observe, that the neighboring cold eddies that are separated by the meandering warm jet in case a), are connected by cold filaments in case b) (e.g. the one in the white rectangle). This implies that the widely used experimental classification of baroclinic waves in a rotating annulus – that is based on the dominant wave number only – is rather incomplete: although the values of Ω and ΔT are within a regime that is (given a certain initialization method, either spin-up or spin-down) characterized by a single dominant wave number m , yet, even *within* this regime, considerable jump-wise state transitions may occur (in terms of pattern topology and also in terms of drift rate) and clearly different dynamical states may develop that essentially have the *same* dominant zonal wave number.

Similarly to the experimental data, a pronounced hysteresis appears at rotation rates $\Omega < 13 \text{ rpm}$ in the cylFloit results (Fig. 10a). In this case the Ω -range of the hysteretic regime clearly agrees with the one found in terms of the dominant wave numbers (cf. Fig. 6b). The interval between the intersection points of the spin-

up and spin-down curves ($\Omega = 6 \text{ rpm}$ and $\Omega = 12 \text{ rpm}$) can therefore be described as the regime where $m = 4$ is the dominant mode of the (lower) spin-down branch *and* the (upper) spin-down branch exhibits $m = 3$. Thus, a manifest correlation is present: at a given Ω the waves of three-fold symmetry propagate faster than the four-fold-symmetric patterns. This conclusion is confirmed by the behavior observed in the from-scratch-initiated simulations of the dark green curve (see also the blue curve of Fig. 6b): in the hysteretic regime, when the system switches from one branch to the other in terms of m , it does so in the drift rate as well. Note, that below $\Omega = 10.1 \text{ rpm}$ (where the aforementioned topological re-organization and sudden drop in the drift rates took place), also in the experimental data of Fig. 10a, the intersection point of the two branches coincides with the onset of the $m = 3$ mode in the spin-up sequence, whereas the spin-down branch maintains the dominant wave number of $m = 4$. In other words: the “first” type of $m = 3$ patterns (seen in Fig. 11a) drifts faster than the baroclinic waves of $m = 4$ at a given rotation rate Ω .

3.4 Empirical Orthogonal Functions

To properly describe the temperature variance stored in co-existent spatio-temporal patterns in the annulus we turned to the method of Empirical Orthogonal Functions

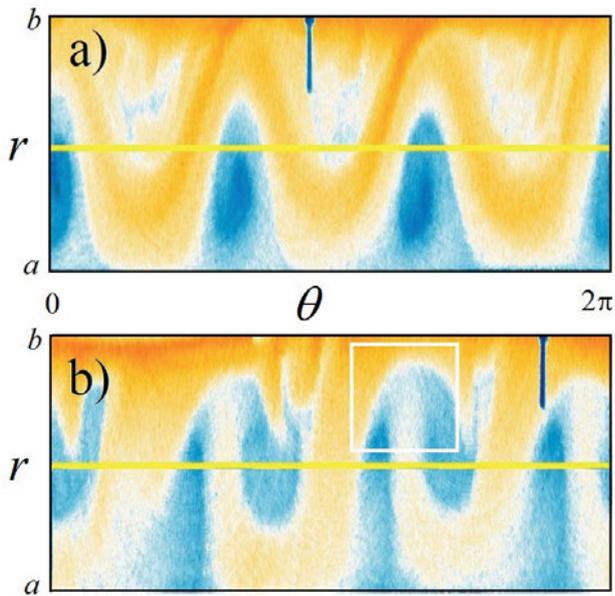


Figure 11: Two thermographic experimental snapshots of $m = 3$ surface temperature patterns. A fastly propagating type (a), observed at rotation rate $\Omega = 4.2$ rpm (see also the corresponding propagation plot in Fig. 9b), and (b) the slower type, observed after the “topological transition” ($\Omega = 10.1$ rpm).

(EOFs) (HARLANDER et al., 2014). This approach is generally accepted as a powerful tool for data compression and dimensionality reduction: it is able to find the spatial patterns of variability, their time variation, and provides a measure for the “relevance” of each pattern, and thus describe the complex behavior of the system, often in terms of surprisingly few modes (VON STORCH and NAVARRA, 1999). It is to be noted, however, that in general these EOF modes do not necessarily correspond to individual dynamical eigenmodes of the system (MONAHAN et al., 2009).

EOF analysis has been extensively used in recent works (HARLANDER et al., 2011; BORCHERT et al., 2014) for two-dimensional temperature and velocity fields in the particular setup at BTU CS. Here, however, as we restricted our studies to the temperature profiles along the circular contour at mid-radius, the one-dimensional EOFs were determined. Organizing the surface temperature data $T(\theta, t)$ at given time instants as column vectors (state vectors) and combining them in temporal order, yields the so-called data matrix \mathbf{X} , whose number of rows and columns correspond to that of the considered spatial and temporal points, respectively. In the present one-dimensional case a transparent visual representation of \mathbf{X}^T can be obtained in the form of a space-time or Hovmöller plot, e.g. the one shown in Fig. 12a (corresponding to an $m = 3$ baroclinic wave).

In our EOF analyses the selected matrices \mathbf{X} consisted of the data from the last 100 time instants of the given (either experimental or numerical) run; a time interval that always lied well within the quasi-stationary part of the investigated process. In space, the experimen-

tal data were linearly interpolated onto an azimuthally equidistant grid of 100 cells, whereas the numerical data were transformed similarly to 50 grid points of uniform spacing. The entries of \mathbf{X} were then obtained by subtracting the mean value of each corresponding row (i.e. temperature time series at a given spatial location). The covariance matrix \mathbf{S} is given by:

$$\mathbf{S} = \frac{1}{n-1} \mathbf{X} \mathbf{X}^T, \quad (3.3)$$

where $n = 100$ is the number of time instants considered. The eigenvectors \mathbf{e}_k (i.e. the EOFs themselves) and the corresponding eigenvalues ξ_k of \mathbf{S} were computed. The EOF index $k = 1, 2, 3, \dots$ is given by organizing the eigenvalues in decreasing order as: $\xi_1 \geq \xi_2 \geq \xi_3 \geq \dots$. The percentage contribution p_k of each pattern \mathbf{e}_k to the total variance captured by the EOFs can then be expressed as: $p_k = \xi_k / \sum_i \xi_i$. As a demonstration, the first four EOF patterns are shown in Fig. 12b, corresponding to the same experiment as the Hovmöller plot of panel a).

3.4.1 Variance distribution

The distribution of percentage contributions p_k of the EOFs (a monotonically decreasing function of index k) was analyzed to quantify the overall complexity of the investigated spatio-temporal patterns. Typical “variability density functions” are presented in Fig. 13a, as obtained from our experiments (black, red and green curves) and the simulations with different models (see also the legend). It is to be emphasized that this figure serves a purely explanatory purpose: to help the reader to better understand the role of the parameters used to quantify the distribution properties. Therefore a large variety of cases at different rotation rates are shown, which are therefore not meant for model comparison. Yet, some common features can be observed: visibly, in a large domain of k , the experimental data points exhibit a power-law type scaling – indicating the importance of higher EOF indices – that is followed by exponential cut-off. A qualitatively similar behavior can be observed in the numerical data as well, however, both the “power-law part” and the “cut-off part” appear to have different quantitative properties than the ones of the experimental results.

To find appropriate measures of these properties, firstly the cumulative density functions $I(k) = \sum_{i=1}^k p_i$ were calculated for each experimental and numerical run. Fig. 13b shows the $I(k)$ curves corresponding to the cases plotted in panel a), with the same color coding. The heuristic empirical form

$$I(k) = 1 - C \frac{e^{-\alpha k}}{k^\beta} \quad (3.4)$$

has proven to be a strikingly accurate parametrization for every run: typically, the asymptotic standard errors were below 3 % for all three free parameters α , β and C .

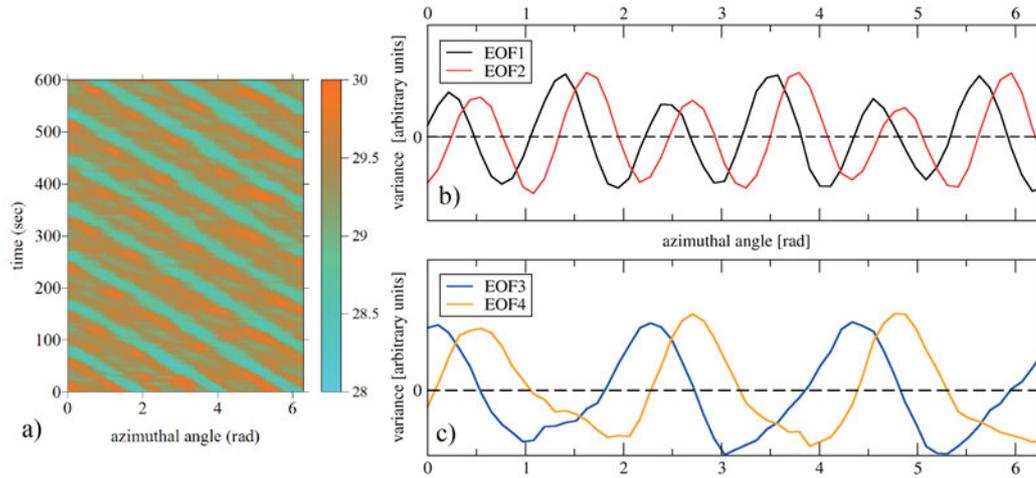


Figure 12: A typical thermographic Hovmöller (space-time) plot of an experimental run at dominant wavenumber $m = 3$ (a), and the first two corresponding EOF variance pattern pairs (b and c). The corresponding relative variances of EOFs 1 to 4 were $p_1 = 0.29$, $p_2 = 0.27$, $p_3 = 0.082$ and $p_4 = 0.073$, respectively.

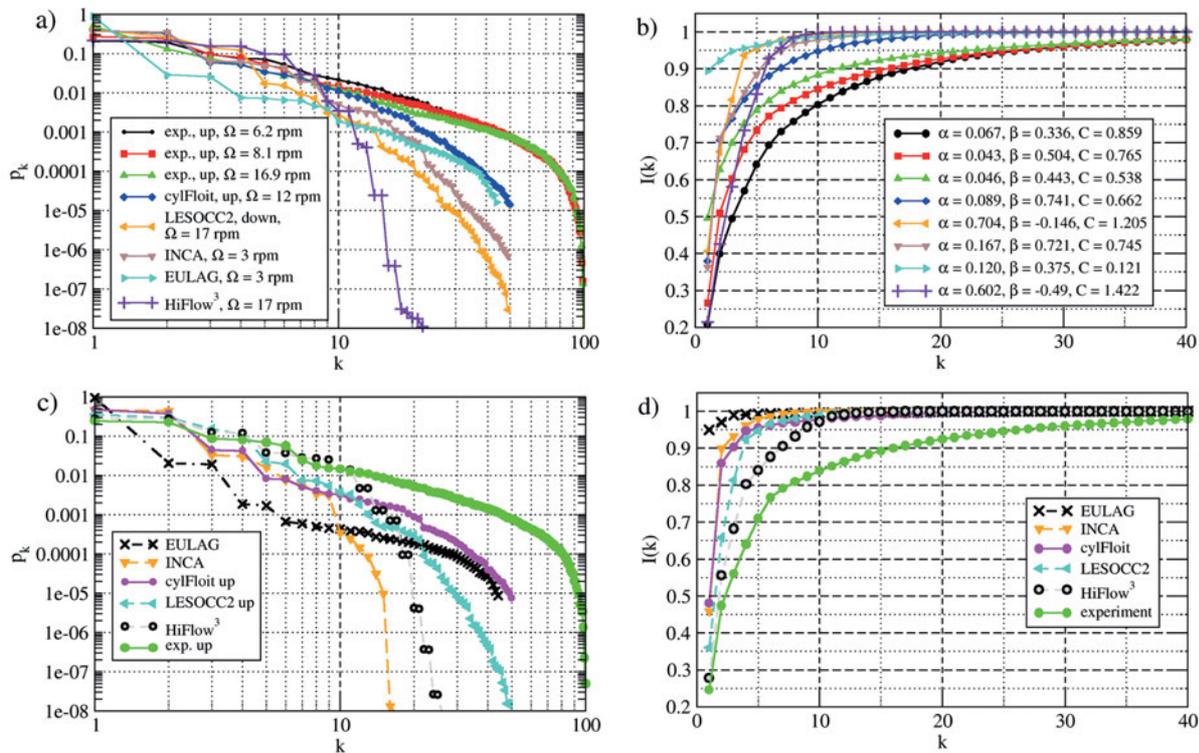


Figure 13: Typical variability density functions obtained from the experiments and numerical models (a). (See legends for the model types and rotation rates). Their corresponding cumulative density functions are shown in panel (b) with the same color coding. The fitted parameter values of α , β and C are also shown. Panels c) and d) show the density functions and cumulative density functions, respectively, for all the models and the experiment for the $\Omega = 9$ rpm (spin-up) case.

Note, that the values of these parameters for the exemplary cases of Fig. 13b are listed in the legend. In panels c) and d) the density functions and cumulative density functions of all the models (and the experiment) are given, all for a single parameter point $\Omega \approx 9$ rpm. For all models, the values of α , β and C were evaluated for each simulated Ω .

Let us now compare the fitted parameters β and α versus rotation rate Ω in Figs. 14a and b, respectively. In the laboratory experiments (red and green curves in both panels) the values of β scatter in the range of $\beta \in (0.3; 1.1)$, while α exhibits small positive values $\alpha \in (0.01; 0.1)$. These imply that the saturation of the cumulative density function is slow, a considerable part

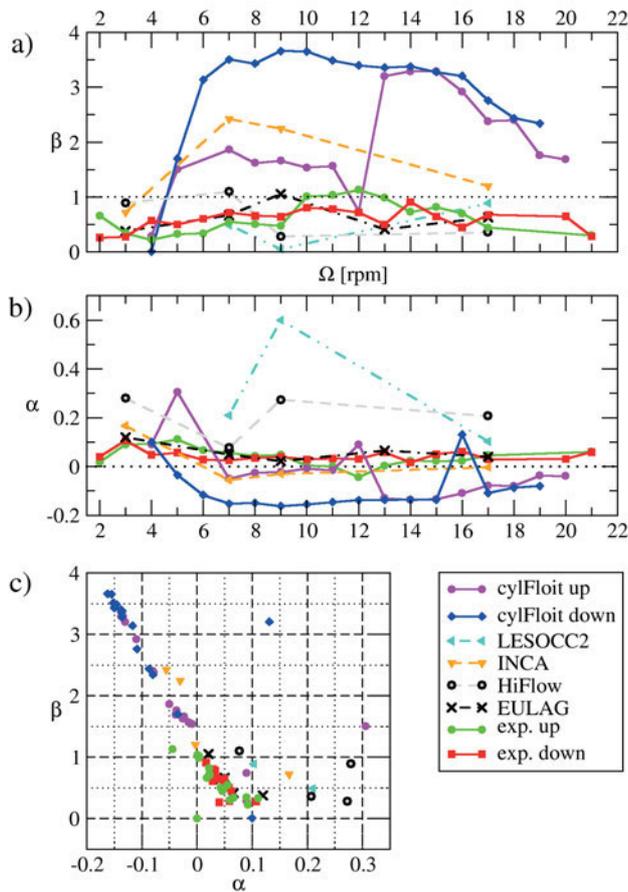


Figure 14: The fitted parameters β and α of the cumulative density functions of (3.4) versus rotation rate Ω : panels a) and b), respectively, and the correlation plot of the two parameters (c). The color coding is the same for all panels.

of the variance is stored in the EOFs of larger k . As the exponential factor is such a slowly varying function (due to the small α), the behavior observed in the experimental density functions of Fig. 13a approximately follows a power-law scaling in the form of $k^{-\gamma} \equiv k^{-\beta-1}$ with $1.3 < \gamma < 2.1$. Such values of γ are typical for the probability density functions of long-range correlated processes. As yet another measure of complexity, it is to be mentioned that $k = 6 - 18$ different EOFs were needed to cover 90 % ($I(k) = 0.9$) of the total variance in the experimental distributions (like the first three graphs listed in Fig. 13b).

The exponent β was also typically found within the same $0 < \beta < 1$ regime in the simulations conducted by EULAG, HiFlow³ and LESOCC2 (see the black, gray and turquoise graphs in Figs. 14a, respectively). This implies that the distribution of variance in these three models behave realistically concerning the smaller k -regime, which practically corresponds to the large-scale features of the flow. Also in terms of α , the EULAG results scattered perfectly within the same interval as the experiments, meaning that the “tail” of the distribution scales correctly. However, the values of parameter C were an order of magnitude smaller for EULAG

($C \in (0.025; 0.12)$) than for all the other cases, either experimental or numerical, where $C \in (0.42; 1.17)$ was found within the baroclinic unstable regime. This is due to the interesting fact that in these simulations – despite of their close-to-perfect scaling properties – the very first EOF alone was responsible for 90–96 % of the total variance, i.e. $p_1 \in (0.9; 0.96)$, a property that can be observed on the turquoise curve of Fig. 13b too. For HiFlow³ and LESOCC2, on the other hand, parameter α appeared to be 2–6 times larger than in the experiments (Fig. 14b), meaning that the variability of larger indices k is suppressed by a marked exponential cut-off, thus most of the variance is stored in the large-scale patterns.

The values of exponent β were found significantly larger in the INCA and cylFloit model runs than in the case of the laboratory experiments (see orange, blue and magenta graphs in Fig. 14a). Typically, the cases where $\beta > 1$ holds, correspond to $\alpha < 0$, as visualized in the correlation plot of Fig. 14c. This relation suggests that at smaller values of index k a sharp “fast” power law characterizes the dominant, large-scale part of the distribution. This scaling, however, is confined only to this regime: in itself it would mean a too sharp cut-off at larger indices k . Thus, for an appropriate parametrization, a negative value of α is needed to compensate this effect to keep the variances at higher EOF indices finite.

Regarding the cylFloit simulations the data points of the spin-up and spin-down series are plotted separately, with magenta and blue symbols in all panels of Fig. 14, respectively. In panel a) the marked hysteresis behavior of parameter β can be observed. This behavior is in manifest connection with the dominant wave numbers (cf. Fig. 6b): apparently, $m = 4$ states are characterized by larger β than $m = 3$ states. This implies that in the $m = 4$ -dominated states the “scale separation” is more pronounced: a larger fraction of the total variance is stored in the first few EOF modes than in the $m = 3$ cases.

In the spin-up and spin-down sequences of the laboratory experiments no such connection was found between wave numbers and the parameters of $I(k)$, however, a significant jump of β at rotation rate $\Omega = 10.1$ rpm is visible in the spin-up curve (green graph in Fig. 14a), that corresponds to the topological transition within the $m = 3$ regime, described in the previous section.

3.4.2 Pattern correlations

Besides the distributions of the eigenvalues of covariance matrix \mathbf{S} , the eigenvectors \mathbf{e}_k , i.e. the variance patterns themselves were also compared. The applied method was similar to the one used in BORCHERT et al. (2014) for two-dimensional EOFs. Firstly, the obtained EOF patterns of indices k and l from the experiment and a given numerical model were linearly interpolated onto the same equidistant grid of 100 cells. These functions are marked by: $f_k^{\text{exp}}(\theta)$ and $f_l^{\text{mod}}(\theta)$, respectively

($\theta \in (0; 2\pi]$). Their correlation coefficient is then calculated as:

$$C_{kl} = \frac{\langle f_k^{\text{exp}}(\theta) f_l^{\text{mod}}(\theta + \varphi) \rangle - \langle f_k^{\text{exp}}(\theta) \rangle \langle f_l^{\text{mod}}(\theta + \varphi) \rangle}{\sigma(f_k^{\text{exp}}(\theta)) \sigma(f_l^{\text{mod}}(\theta + \varphi))}, \quad (3.5)$$

where $\langle \cdot \rangle$ marks the azimuthal mean, $\sigma(\cdot)$ denotes the standard deviation and φ is the “offset angle” which maximizes C_{kl} . This sliding transformation is required, since the azimuthal orientation of EOFs in the various models (and experimental runs) are generally different. In this transformation periodic boundary conditions were applied, i.e. the values for which $\theta + \varphi > 2\pi$ were actually mapped onto the interval $(0; \varphi)$.

The values C_{kl} were calculated for the first 10 EOFs (both numerical and experimental) and were combined into 10×10 matrices. The structures of these matrices were analyzed. Here, we present a few typical exemplary cases to yield a qualitative insight to the nature of the correlation properties of one-dimensional EOFs. In Fig. 15 the correlation plots for the benchmark case #4 ($\Omega \approx 17$ rpm) are presented. This case was selected, since here – already out of the hysteretic regime – all the models found $m = 4$ as dominant mode, in agreement with the experiments. For a better understanding of the comparisons to follow, in panel a) we present the correlation plot of the EOFs of the given experiment with one another (hence, $f_i^{\text{exp}} \equiv f_i^{\text{mod}}$ using the above notation). Trivially, in this case $C_{ii} = 1$ holds for the diagonal entries, and the matrix is symmetric. Though the EOFs are, by definition, orthogonal, yet, the aforementioned sliding transformation leads to rather marked correlations, since EOF1 and EOF2 (and, similarly EOF3 and EOF4, etc.) are rather similar, but shifted in azimuthal direction (see also Fig. 12b and c). Such EOF pairs account for the baroclinic wave propagation, analogously to the relation of sine and cosine terms in the Fourier decomposition of propagating waves.

Panels b)–f) show the correlation matrices obtained from the comparison of the experimental set of EOFs with the EOFs from cylFloIt, EULAG, HiFlow³, INCA and LESOCC2, respectively. The numbers on the horizontal axis represent the indices of the experimental variance patterns, and those on the vertical axis are the EOF indices of the given numerical model. The indices and values of the maximum entries in the given matrix are also marked in the panels. Two main observations need to be emphasized. Firstly, the structures of the matrices are rather similar, implying that the numerical models produce similar variance patterns. Also, the aforementioned EOF pairs are clearly visible in the matrices in the form of 2×1 and 2×2 blocks of closely similar correlations. The second main observation is that, despite of the similarity of the matrices, none of them has diagonal structure as in panel a). Thus, the various EOF patterns are ranked differently.

The latter statement is seemingly in contrast with the findings of BORCHERT et al. (2014), who found correlation coefficients above 0.9 by comparing their EOFs (ob-

tained using the cylFloIt and INCA codes) to the laboratory EOFs of the *same* index. However, there the full two-dimensional surface temperature patterns were taken. As a test of consistency, we applied our methodology to the very same experimental records from year 2011 and the same (“from-scratch” initiated) cylFloIt runs studied in BORCHERT et al. (2014) to obtain the correlation coefficients for the one-dimensional EOFs. The resulting correlation matrix is shown in Fig. 16a. Apparently, the obtained structure is quite similar to those seen in Fig. 15b–f, and lacks large values in the diagonal. However, the entries in the 2×2 blocks in the vicinity of the diagonal at lower left are indeed large, with a maximum of $C_{31} = 0.97$. The similarities and differences of these patterns can be visually evaluated in Fig. 16b and c, where EOFs 1 and 3 are plotted for the experimental and the numerical case, respectively. One can see, that in the experimental case EOF1 exhibits wave number $m = 6$ (and so does its shifted pair EOF2, not shown here) and the dominant baroclinic wave number $m = 3$ appears in the EOF3 for the first time, in contrast to the typical numerical results. Thus, the numerical models have a tendency to underestimate the variance stored in the smaller scales.

It can be stated that the one-dimensional data extracted from the surface temperature field at mid-radius r_{mid} are generally more sensitive to smaller-scale differences than the full two-dimensional patterns, since – as discussed above – in the two-dimensional case no such “EOF swap” occurs between numerics and experiment. The mid-radius temperature profiles are apparently largely effected by the variance stored in the harmonics of the dominant baroclinic wave mode, related to the structure and dynamics of the cold eddies in the lobes of baroclinic waves. The fact that the numerical models are apparently not able to resolve these phenomena implies that they may well be related to boundary layer effects or even “wind” stress above the free surface of the laboratory tank, which are clearly out of the scope of the investigated numerical models.

Also, it is to be noted, that in an annulus with an exact rotational invariance the EOFs must be sinusoidal, i.e. each would project on a single azimuthal wave number only, as shown by ACHATZ and SCHMITZ (1997). The fact that the typical EOFs of the experiment can in many cases visibly be decomposed to at least two wave numbers (as the ones in Fig. 12b and Fig. 16b) indicates a violation of rotational symmetry and nonlinear dynamics. In the azimuthally invariant numerical models (as cylFloIt), however, the EOF patterns were indeed found to be nearly sinusoidal (see e.g. Fig. 16c). Their slight imperfection is merely a consequence of the finite length of the time series considered.

4 Summary and conclusions

In this work we have critically compared various experimentally and numerically obtained characteristic properties of baroclinic instability in a differentially heated

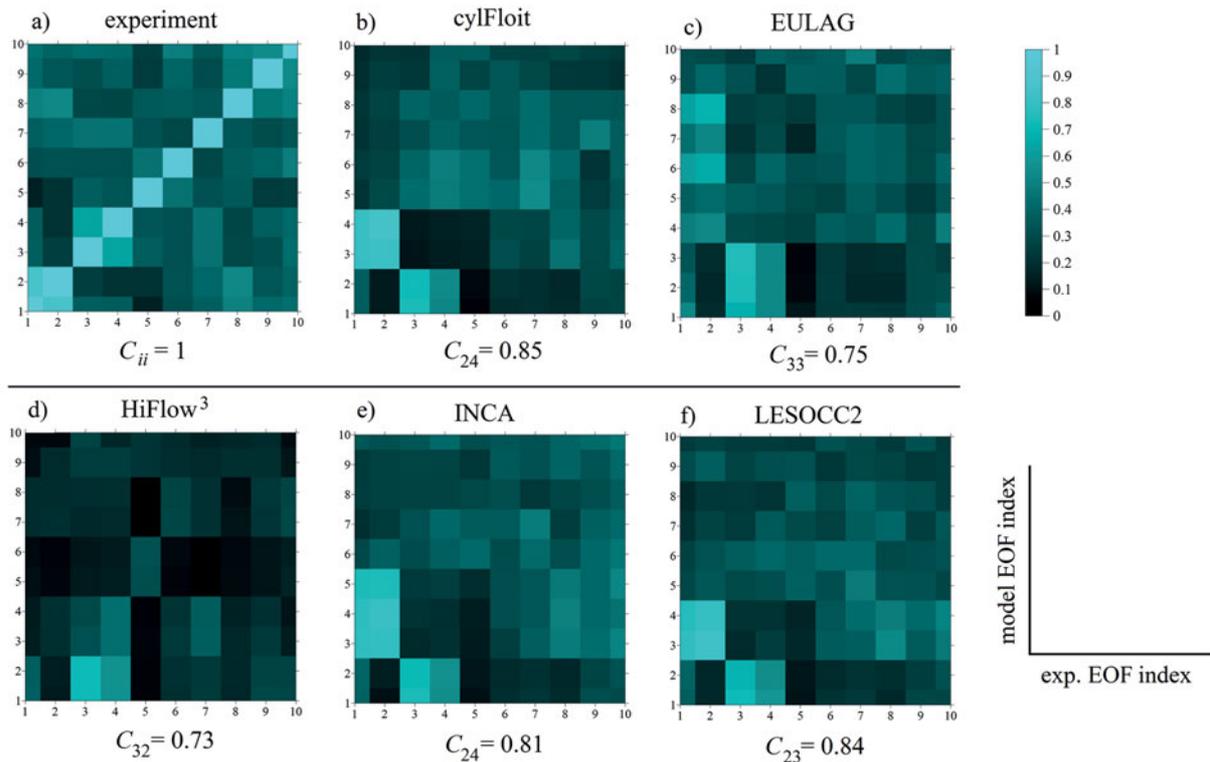


Figure 15: The cross-correlation matrices obtained in the benchmark case #4 ($\Omega \approx 17$ rpm). The positions and values of the maximum entries of the matrices are also given underneath the respective figures.

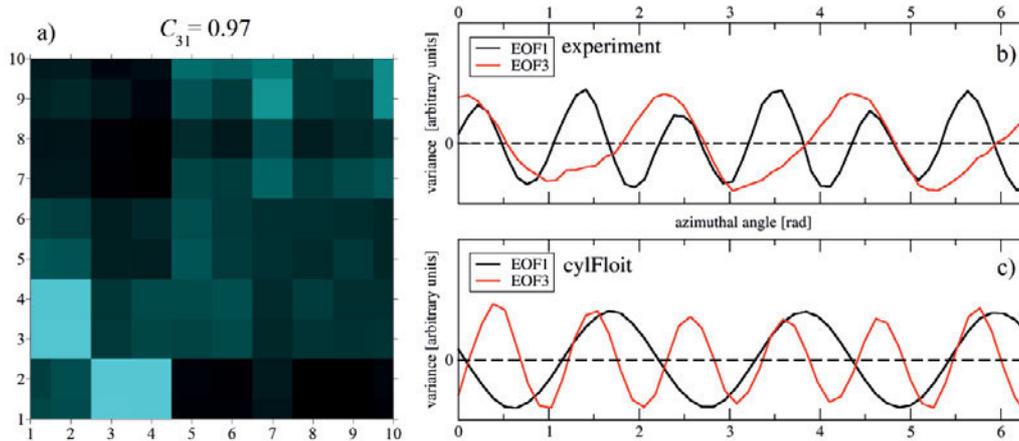


Figure 16: The correlation matrix of the one-dimensional EOFs, obtained from the numerical and experimental data of test case #7 of BORCHERT et al. (2014), and the value of the maximum entry (a). EOFs 1 and 3 of the experimental (a) and numerical (b) case. Note the “swap” between the indices and wave patterns of the two cases.

rotating annulus. Our systematic comparison of five different numerical models to laboratory experiments (“benchmarking”) was largely motivated by the general need to validate numerical models and procedures to be used for modeling large-scale atmospheric flows.

Two series of laboratory measurements were performed: the “spin-up” and “spin-down” sequences. Between each measurement only rotation rate Ω was adjusted, while the radial temperature difference $\Delta T \approx 8$ K remained constant. The two sequences enabled us

to scan through the investigated parameter range with different initial conditions, and thus access multiple equilibrium regimes. In agreement with earlier results (MILLER and BUTLER, 1991; SITTE and EGBERS, 2000; VON LARCHER et al., 2005) a considerable hysteresis was found in terms of the dominant azimuthal wave numbers m of the baroclinic waves.

It is well established since the works of JAMES et al. (1981) and HIGNETT et al. (1985) in the 1980s, that in terms of m , the development of baroclinic waves in baro-

clinic annuli can be captured in direct numerical simulations fairly well. In the present work we also found that m is indeed a robust indicator of the flow state, and its obtained values exhibit good agreement between the experiments and the numerical runs. The numerical results also support our conclusion that the hysteretic behavior of m is to be interpreted as distinct multiple equilibria and is not just caused by transient phenomena. This statement is backed by the following observations: (i) Simulation series conducted with models cylFloit and LESOCC2 imitated the “spin-up” and “spin-down” sequences and found hysteresis in terms of m . A third bunch of simulations, however, were always initialized from the axially symmetric stable state (cylFloit “from scratch” sequence). Yet, occasionally even here, wave numbers characteristic to the “spin-down” branch were found to develop within a rotation rate regime where these simulations typically converged to the states of the “spin-up” branch. (ii) In the HiFlow³ simulations, runs with slightly perturbed initial conditions were also conducted. The only cases where these temperature disturbances yielded a different dominant wave number m than that corresponding to the unperturbed runs were at parameter points within the experimentally observed hysteretic Ω -regime.

The different routes of releasing the stored potential energy – as initial conditions – may play important roles in the observed hysteretic mode (i.e. dominant m -) selection. If spin-up is applied, the inclination of the iso-density surfaces increases, making the flow more “baroclinic unstable”, i.e. the stored potential energy is transformed into kinetic energy only via enhanced baroclinic wave excitation (a very effective way of energy release). If spin-down is applied, the inclination of isopycnals decrease, and the unavoidable transient hydrostatic imbalance opens up the way to other forms of kinetic energy release as well, via the overturning background flow (possibly even inertia-gravity wave excitation, see e.g. the numerical work of RANDRIAMAMPINANINA et al. (2006)). Although these effects are probably transient, yet, as initial conditions they appear to play an important role in setting the dominant wave number m of the baroclinic wave, which remains steady for a much longer time than the timescale characteristic for the reorganization of stratification.

Another important measure of baroclinic wave dynamics is the drift rate c_m of the dominant wave mode. In qualitative agreement with the quasigeostrophic Eady model (VALLIS, 2006), the $c_m(\Omega)$ relationship was found to be a decreasing function, roughly following the $c_m \propto \Omega^{-1}$ dependence set by the thermal wind balance. It is to be noted, however, that most of the models (with the exception of EULAG) systematically overestimated the wave speeds. This phenomenon may well be explained by the simulations’ difficulties to resolve the boundary layer drag at the lateral sidewalls. A similar observation was described in the study of WILLIAMS et al. (2010) where a two-layer (lid shear-driven) rotating baroclinic annulus set-up was investigated both ex-

perimentally and numerically. In their case the simulated drift rates were larger than the measured values by a factor of 4, due to the model’s neglect of Stewartson layer drag. Stewartson layers are characteristic for homogeneous fluids. In our case of relatively strong stratification, however, $PrRo_T/\Gamma^2 \gg Ek^{2/3}$ holds with Γ being the vertical aspect ratio of the tank (as defined in Section 2) and $Ek = \nu/(\Omega L^2)$ the Ekman number. In this regime – instead of Stewartson layers – two boundary layers are found in the vicinity of each lateral sidewall: the larger hydrostatic layer with a characteristic thickness of $\delta_h = D(PrRo_T/\Gamma^2)^{1/2}$ and, closer to the wall, the buoyancy layer whose thickness is $\delta_b = D(\nu\kappa/(D^3 g\alpha\Delta T))^{1/4}$. These two layers unite and form the Stewartson layer (with $\delta_s = DEk^{1/3}$) if stratification decreases (BARCILON and PEDLOSKY, 1967). For the present case $\delta_h > b - a$ holds, i.e. practically the whole measurement cavity lies within the “hydrostatic” domain. The buoyancy layer, however, is found to be only $\delta_b \approx 1$ mm thick, thus it is not resolved sufficiently by most of the models.

The sensitivity of drift rates to the horizontal grid spacing was demonstrated with the INCA model. The phase speeds of baroclinic waves at two different rotation rates – namely $\Omega = 4$ rpm and $\Omega = 9.5$ rpm – were determined using two grids in both cases for comparison. The coarse and fine grids had minimum cell sizes of $\Delta x_{\min,(1)} = \Delta y_{\min,(1)} = 1.5$ mm and $\Delta x_{\min,(2)} = \Delta y_{\min,(2)} = 0.5$ mm, respectively. The obtained drift rates were: $c^{(1)} = 0.097$ rad/s (coarse grid); $c^{(2)} = 0.057$ rad/s (fine grid) at $\Omega = 4$ rpm, and $c^{(1)} = 0.025$ rad/s (coarse grid); $c^{(2)} = 0.023$ rad/s (fine grid) at $\Omega = 9.5$ rpm. Visibly, at the lower rotation rate (where the phase velocities of baroclinic waves are generally large) the refinement of the horizontal grid yielded slower wave propagation almost by a factor of two. In the case of the higher rotation rate this effect was manifestly smaller – around 10% – in qualitative agreement with the drag-hypothesis: the drag itself is expected to be smaller too if the drift itself is slower. Thus, we can conclude that the grid resolution has marked effect on the simulated wave speeds, and to get a proper insight into the flow structure at the vicinity of the lateral sidewalls, one needs to apply grids that properly resolve the buoyancy layer.

We also found marked connection between the spatial patterns of baroclinic waves and their drift rates, both experimentally and numerically. The aforementioned hysteresis that was observed in terms of the dominant wave number m also manifested itself in the drift rates. In the cylFloit simulations, $m = 3$ waves always propagated faster than their $m = 4$ counterparts at a given rotation rate (within the hysteretic Ω -regime). Similar behavior was noticed in the laboratory experiments too: a certain type of the $m = 3$ wave was found to be faster than the $m = 4$ waves of the same Ω . However, in the laboratory, another type of three-fold symmetric ($m = 3$) pattern appeared as well in the “spin-up” series, which was found to propagate at an even smaller speed

than the $m = 4$ waves. Here the surface temperature pattern has undergone a “topological” reorganization: the meandering warm jet that separated the inner and outer domain in the “fast” $m = 3$ waves has disconnected. This transition possibly opens the way for stronger radial temperature fluxes, therefore this new configuration may reduce the thermal wind (background flow) more effectively, yielding slower drift. Applying the same reasoning for the hysteresis of $m = 4$ waves and the “fast” $m = 3$ waves, it can be stated that among these, the $m = 4$ mode exhibits larger radial heat flow. As far as the general heat flow is considered, RAYER et al. (1998) showed that the Nusselt number Nu in a baroclinic annulus exhibits a large drop at the transition from axisymmetric flow to the regime of regular waves, where – compared to the abrupt change at the onset of baroclinic instability – it does not change markedly with the increasing Ω . This plateau of $Nu(\Omega)$ is followed by another pronounced drop of Nu when the system reaches higher rotation rates where the waves become irregular (this state was not studied in the present work). It is to be noted, that the changes in heat flow that can be attributed to shape changes of regular baroclinic waves is rather small compared to the aforementioned changes between the general flow states. We also remark, that – as demonstrated in the experiments of FEIN (1973) – the drift rates are also highly sensitive to the upper boundary condition, that was not prescribed properly in the model equations.

The third main focus of our study was the statistical quantification of the structures of the surface temperature field and the analysis of their spatio-temporal variability. As a measure of the overall variability in the system, the time averaged temperature variance $\bar{\sigma}(\Omega)$ taken along the circular contour at mid radius r_{mid} served as an order parameter that indicates the breaking of the axial symmetry (and, thus, the onset of baroclinic instability) with a marked jump at critical rotation rate Ω_{crit} . Indeed, in the numerical simulations $\bar{\sigma} \approx 0$ was detected in all cases where no dominant wave mode could be found (aside for the trivial $m = 0$), implying the stability of the axially symmetric basic state. This was then followed by more than 10 times larger variances at $\Omega > \Omega_{\text{crit}}$. However, in the laboratory experiments the transition was not that apparent: even below Ω_{crit} fluctuations appeared on the same order of magnitude as the $\bar{\sigma}$ values of higher rotation rates (though, smaller by a factor of around 0.5). This observation confirms our previous finding of spontaneous excitation of dispersive transient wave-like phenomena (coined “weak waves”) that “blur” the boundary of instability in the parameter space (VINCZE et al., 2014). This qualitative difference between numerics and experiments indicates the presence of non-modal transient growth of small temperature fluctuations in this sensitive regime (SEELIG et al., 2012) unavoidable in the laboratory (see also the work of HOFF et al. (2014) in the present volume). In the numerical results the temperature variance obtained at a few centimeters below the surface was found to be significantly larger (by a

factor of around 2) than at the surface. This behavior, however, could not be verified experimentally with the applied measurement techniques.

In order to analyze smaller scale spatial structures, we calculated the Fourier spectra of the azimuthal temperature profiles along the circular contour at mid-radius r_{mid} for all time instants of a given experimental or numerical run, and their temporal average was considered as the characteristic spectral “fingerprint” of the investigated pattern. In the case of an m -fold symmetric baroclinic wave, besides the dominant wave number, its harmonics also appear in the spectra with finite amplitudes, as already demonstrated by JAMES et al. (1981). The amplitudes and the significance of the spectral peaks provide a measure of the importance of the regular smaller scale patterns. Typically, in the experimental data marked amplitudes were observed at the integer multiples of the dominant mode. These amplitudes were, in many cases, comparably large to that of the dominant wave number (representing the overall rotational symmetry). In the cylFloit simulations however, the harmonics were not that pronounced. These smaller-scale patterns are attributed to the cold eddies outside and inside the meandering jet of the baroclinic wave. The fact that these structures could not be resolved accurately in the simulations may be due to the following reasons: the cold eddies in the vicinity of the outer rim seem to be excited by shear instability involving the buoyancy layer, which was not resolved by most of the models, as discussed above. This statement is supported by the observation that the high-resolution EULAG temperature fields occasionally showed such structures (cf. VON LARCHER and DÖRNBRACK (2014) in the present issue). Surface phenomena that are out of the scope of the studied governing equations may also contribute to making these cold eddies more pronounced. Such factors can be e.g. the “wind” stress that takes place at the free surface of the experimental tank as it rotates, or the presence of finite vertical heat fluxes at the top surface (note, that all the models included the $\nabla T \cdot \mathbf{e}_z|_{z=D} = 0$ type no-flux boundary conditions, which certainly cannot be achieved in the experiment due to the free surface). The role of the slight curvature of the top surface due to the centrifugal effect is probably negligible: the water depth difference Δd between the inner and outer rim is found to be $\Delta d = \Omega^2(b^2 - a^2)/(2g) \approx 0.4 - 2.5$ mm in the applied rotation rate regime. Nevertheless, this minor geometric distortion also adds to the list of effects that were not implemented in the numerical models.

The azimuthal temperature variance patterns were decomposed into sets of empirical orthogonal functions (EOFs). We found that in the experimental distribution of the ranked relative variances – the normalized eigenvalues corresponding to the EOF modes – typically follows a slowly decaying power-law type scaling, implying that a considerable part of the total variance is stored in the smaller scales (6–18 orthogonal modes were needed to cover 90 % of the total variance). In general, the numerically obtained distributions exhibited

faster cut-offs towards the higher ranks, thus less small-scale variance. The practical absence of the correlated small-scale thermal fluctuations in the simulations supports the need for some subgrid-scale parametrization that takes into account the growth of temperature fluctuations that might play a significant role in the dynamics. These fluctuations can be caused by the aforementioned experimental impurities (or possibly induced by boundary layer effects) and “inflated” through the nonlinear interactions. Since the present work (and the MetStröm collaboration in general) has been motivated by meteorological problems, it is important to emphasize, that obviously, the sub-grid-scale processes of the atmosphere are very different from, and way more complex than, those of our experimental set-up. Yet, our observations of how the small-scale variability affects the stability of the larger-scale patterns in the set-up – regardless of the actual physical processes at subgrid-scales – from a more general, dynamical systems point of view, may also be of relevance for atmospheric or oceanic modeling.

Due to the differences between the numerical procedures, drawing general summarizing conclusions from the present study is far from straightforward. The codes built by the numerical groups of our collaboration were all developed based on the already existing solvers and numerical schemes that the respective groups have previously gained experience with, independently from modeling the differentially heated rotating annulus. Therefore, since the models were different in so many aspects (reference frames, grid structures, resolutions, subgrid-scale parametrizations, etc.), the “parameter space” of model features was so large, that no benchmarking strategy could have been intentionally defined for comparing models of increasing degrees of complexity from one single aspect.

However, the main message of our results is that despite of all this large diversity of the applied codes, they do share some interesting commonalities. Firstly, the experimentally obtained dominant wave numbers were fairly well reproduced by all the codes. This is somewhat surprising in the light of the fact that the drift rates (that are closely related to strength of the zonal background flow) were systematically overestimated by all the models, except EULAG. This latter common feature was found to be the consequence of improper grid resolution at the vicinity of the vertical sidewalls, where the boundary layer drag was not sufficiently resolved. As another general “message to take home” we can also state that (i) the small-scale variability was not captured properly by the models, as discussed above, yet (ii), the large-scale patterns are properly reproduced. Thus, despite of the obvious nonlinearity of the studied dynamics, we can state that the coupling (‘information transfer’) between the smaller and larger scales does not play such a critical role in the studied flow regimes.

As a possible extension and continuation of this idea, the response of the system to small amplitude temporal and spatial thermal fluctuations (entering via the bound-

ary conditions) could be analyzed numerically in a future research project. Such investigations – if the above assumptions are correct – can possibly lead to even more accurate numerical modeling and a deeper understanding of the dynamics in this set-up. Also, our future plans involve the extension of the presented benchmarking techniques to numerical methods that reach beyond the Boussinesq approximation (e.g. Low-Mach models) whose application may be wise in the larger ΔT -regime.

The results presented in this paper have clearly demonstrated that the relatively simple rotating annulus arrangement indeed provides a remarkable test bed to verify and tune numerical methods aiming to model large-scale atmospheric flows. The authors think that the presented pool of experimental and numerical data and the applied evaluation methods and “test quantities” will also prove useful benchmarks for similar studies in the future.

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References

- ACHATZ, U., G. SCHMITZ, 1997: On the closure problem in the reduction of complex atmospheric models by PIPs and EOFs: A comparison for the case of a two-layer model with zonally symmetric forcing. – *J. Atmos. Sci.* **54**, 2452–2474.

- BARCILON, V., J. PEDLOSKY, 1967: A unified linear theory of homogeneous and stratified rotating fluids. – *J. Fluid Mech.* **29**, 609–621.
- BORCHERT, S., U. ACHATZ, S. REMMLER, S. HICKEL, U. HARLANDER, M. VINCZE, K.D. ALEXANDROV, F. RIEPER, T. HEPPELMANN, S.I. DOLAPCHIEV, 2014: Finite-volume models with implicit subgrid-scale parameterization for the differentially heated rotating annulus. – *Meteorol. Z.* **23**, this issue.
- BREZZI, F., 1974: On the existence, uniqueness, and approximation of saddle point problems arising from Lagrangian multipliers. – *Revue française d'automatique, informatique, recherche opérationnelle. Analyse Numérique* **8**, 129–151.
- CHORIN, A.J., 1968: Numerical solution of the Navier–Stokes equations. – *Math. Comp.* **22**, 745–762.
- COTTER, C., P. SMOLARKIEWICZ, I. SZCZYRBA, 2002: A viscoelastic fluid model for brain injuries. – *Int. J. Num. Meth. Fluids* **40**, 303–311.
- EADY, E., 1949: Long waves and cyclone waves. – *Tellus* **1**, 33–52.
- ELLIOTT, J., P. SMOLARKIEWICZ, 2002: Eddy resolving simulations of turbulent solar convection. – *Int. J. Num. Meth. Fluids* **39**, 855–864.
- FEIN, J.S., 1973: An experimental study of the effects of the upper boundary condition on the thermal convection in a rotating, differentially heated cylindrical annulus of water. – *Geophys. Astrophys. Fluid Dyn.* **5**, 213–248.
- FRÖHLICH, J., 2006: Large Eddy Simulation turbulenter Strömungen – Teubner Verlag, 414 pp.
- FRÜH, W.-G., P. READ, 1997: Wave interactions and the transition to chaos of baroclinic waves in a thermally driven rotating annulus. – *Philosophical Transactions of the Royal Society of London. Series A: Math. Phys. Engin. Sci.* **355**, 101–153.
- FULTZ, D., R.R. LONG, G.V. OWENS, J. WEIL, 1959: Studies of thermal convection in a rotating cylinder with some implications for large-scale atmospheric motions – American Meteorological Society, Meteorological monographs, v. 4, no. **21**.
- GOLDSTEIN, D., R. HANDLER, L. SIROVICH, 1993: Modeling a no-slip flow boundary with an external force field. – *J. Comput. Phys.* **105**, 354–366.
- GRABOWSKI, W., P. SMOLARKIEWICZ, 2002: A multiscale anelastic model for meteorological research. – *Mon. Wea. Rev.* **130**, 939–956.
- GRILLI, M., P.J. SCHMID, S. HICKEL, N.A. ADAMS, 2012: Analysis of unsteady behaviour in shockwave turbulent boundary layer interaction. – *J. Fluid Mech.* **700**, 16–28.
- GYÜRE, B., I. BARTOS, I.M. JÁNOSI, 2007: Nonlinear statistics of daily temperature fluctuations reproduced in a laboratory experiment. – *Phys. Rev. E* **76**, 037301.
- HARLANDER, U., R. FAULWETTER, K. ALEXANDROV, C. EGBERS, 2009: Estimating local instabilities for irregular flows in the differentially heated rotating annulus. – In: *Advances in Turbulence XII*, Springer, 163–166.
- HARLANDER, U., T. VON LARCHER, Y. WANG, C. EGBERS, 2011: PIV and LDV measurements of baroclinic wave interactions in a thermally driven rotating annulus. – *Exp. Fluids* **51**, 37–49.
- HARLANDER, U., J. WENZEL, K. ALEXANDROV, Y. WANG, C. EGBERS, 2012: Simultaneous PIV and thermography measurements of partially blocked flow in a differentially heated rotating annulus. – *Exp. Fluids* **52**, 1077–1087.
- HARLANDER, U., T. VON LARCHER, G.B. WRIGHT, M. HOFF, K.D. ALEXANDROV, C. EGBERS, 2014: Orthogonal decomposition methods to analyze PIV, LDV and thermography data of a thermally driven rotating annulus laboratory experiment. – In: T. von Larcher, P.D. Williams (eds.): *Modelling Atmospheric and Oceanic flows: Insight from Laboratory Experiments and Numerical Simulations*. Wiley, DOI: [10.1002/9781118856024.ch17](https://doi.org/10.1002/9781118856024.ch17)
- HEUVELINE, V., 2010: HiFlow3: A flexible and hardware-aware parallel finite element package. – In: *Proceedings of the 9th Workshop on Parallel/High-Performance Object-Oriented Scientific Computing, POOSC '10*, 4:1–4:6, New York, NY, USA. ACM.
- HEUVELINE, V., E. KETELAER, S. RONNAS, M. SCHMIDT-BREICK, M. WLOTZKA, 2012: Scalability study of HiFlow3 based on a fluid flow channel benchmark. – *EMCL Preprint Series* (2012/1).
- HICKEL, S., N.A. ADAMS, J.A. DOMARADZKI, 2006: An adaptive local deconvolution method for implicit LES. – *J. Comput. Phys.* **213**, 413–436.
- HICKEL, S., T. KEMPE, N.A. ADAMS, 2008: Implicit large-eddy simulation applied to turbulent channel flow with periodic constrictions. – *Theo. Comput. Fluid Dyn.* **22**, 227–242.
- HIGNETT, B.P., A. WHITE, R. CARTER, W. JACKSON, R. SMALL, 1985: A comparison of laboratory measurements and numerical simulations of baroclinic wave flows in a rotating cylindrical annulus. – *Quart. J. Roy. Meteor. Soc.* **111**, 131–154.
- HINTERBERGER, C., J. FRÖHLICH, W. RODI, 2007: Three-dimensional and depth-averaged large-eddy simulations of some shallow water flows. – *J. Hydraulic Engin.* **133**, 857–872.
- HOFF, M., U. HARLANDER, C. EGBERS, 2014: Empirical singular vectors of baroclinic flows deduced from experimental data of a differentially heated rotating annulus. – *Meteorol. Z.* **23**, 581–597, DOI: [10.1127/metz/2014/0563](https://doi.org/10.1127/metz/2014/0563)
- JAMES, I., P. JONAS, L. FARNELL, 1981: A combined laboratory and numerical study of fully developed steady baroclinic waves in a cylindrical annulus. – *Quart. J. Roy. Meteor. Soc.* **107**, 51–78.
- JÁNOSI, I.M., P. KISS, V. HOMONNAI, M. PATTANTYÚS-ÁBRAHÁM, B. GYÜRE, T. TÉL, 2010: Dynamics of passive tracers in the atmosphere: Laboratory experiments and numerical tests with reanalysis wind fields. – *Phys. Rev. E* **82**, 046308.
- LORENZ, E.N., 1963: The mechanics of vacillation. – *J. Atmos. Sci.* **20**, 448–465.
- MASON, P., 1975: Baroclinic waves in a container with sloping end walls. – *Philosophical Transactions of the Royal Society of London. Series A, Math. Phys. Sci.* **278**, 397–445.
- MAYER, J., 2007: Ilu++: A new software package for solving sparse linear systems with iterative methods. – *PAMM* **7**, 2020123–2020124.
- MEYER, M., S. HICKEL, N. ADAMS, 2010: Assessment of implicit large-eddy simulation with a conservative immersed interface method for turbulent cylinder flow. – *Int. J. Heat Fluid Flow* **31**, 368–377.
- MILLER, T.L., K.A. BUTLER, 1991: Hysteresis and the transition between axisymmetric flow and wave flow in the baroclinic annulus. – *J. Atmos. Sci.* **48**, 811–824.
- MONAHAN, A.H., J.C. FYFE, M.H. AMBAUM, D.B. STEPHENSON, G.R. NORTH, 2009: Empirical orthogonal functions: The medium is the message. – *J. Climate* **22**, 6501–6514.
- PFEFFER, R.L., W.W. FOWLIS, 1968: Wave dispersion in a rotating, differentially heated cylindrical annulus of fluid. – *J. Atmos. Sci.* **25**, 361–371.
- PRUSA, J., P. SMOLARKIEWICZ, A. WYSZOGRODZKI, 2008: EU-LAG, a computational model for multiscale flows. – *Computers & Fluids* **37**, 1193–1207.
- RANDRIAMAMPANINA, A., W.-G. FRÜH, P.L. READ, P. MAUBERT, 2006: Direct numerical simulations of bifurcations in an air-filled rotating baroclinic annulus. – *J. Fluid Mech.* **561**, 359–389.
- RAVELA, S., J. MARSHALL, C. HILL, A. WONG, S. STRANSKY, 2010: A realtime observatory for laboratory simulation of planetary flows. – *Experiments in Fluids* **48**, 915–925.

- RAYER, Q., D. JOHNSON, R. HIDE, 1998: Thermal convection in a rotating fluid annulus blocked by a radial barrier. – *Geophys. Astrophys. Fluid Dyn.* **87**(3-4), 215–252.
- READ, P.L., 2003: A combined laboratory and numerical study of heat transport by baroclinic eddies and axisymmetric flows. – *J. Fluid Mech.* **489**, 301–323.
- READ, P., S. LEWIS, R. HIDE, 1997: Laboratory and numerical studies of baroclinic waves in an internally heated rotating fluid annulus: a case of wave/vortex duality?. – *J. Fluid Mech.* **337**, 155–191.
- READ, P.L., P. MAUBERT, A. RANDRIAMAMPANINA, W.-G. FRÜH, 2008: Direct numerical simulation of transitions towards structural vacillation in an air-filled, rotating, baroclinic annulus. – *Phys. Fluids (1994-present)* **20**, 044107.
- REMMLER, S., S. HICKEL, 2012: Direct and large eddy simulation of stratified turbulence. – *Int. J. Heat Fluid Flow* **35**, 13–24.
- REMMLER, S., S. HICKEL, 2013: Spectral structure of stratified turbulence: Direct numerical simulations and predictions by large eddy simulation. – *Theo. Comput. Fluid Dyn.* **27**, 319–336.
- RHIE, C.M., W.L. CHOW, 1983: Numerical study of the turbulent flow past an airfoil with trailing edge separation. – *AIAA J.* **21**, 1525–1532.
- SCHRÖTTLE, J., A. DÖRNBRACK, 2013: Turbulence structure in a diabatically heated forest canopy composed of fractal pythagoras trees. – *Theo. Comput. Fluid Dyn.* **27**, 337–359.
- SEELIG, T., U. HARLANDER, R. FAULWETTER, C. EGBERS, 2012: Irregularity and singular vector growth in the differentially heated rotating annulus. – *Theo. Comp. Fluid Dyn.* **27**, 415–432.
- SHU, C.-W., 1988: Total-variation-diminishing time discretizations. – *SIAM Journal of Scientific and Statistical Computing* **9**, 1073–1084.
- SITTE, B., C. EGBERS, 2000: Higher order dynamics of baroclinic waves. – In: *Physics of Rotating Fluids*, Springer, 355–375.
- SMOLARKIEWICZ, P., 1991: On forward-in-time differencing for fluids. – *Monthly Weather Review* **119**, 2505–2510.
- SMOLARKIEWICZ, P., L. MARGOLIN, 1997: On forward-in-time differencing for fluids: An eulerian/semi-lagrangian nonhydrostatic model for stratified flows. – *Atmos.-Ocean Special* **35**, 127–157.
- SMOLARKIEWICZ, P., L. MARGOLIN, 1998: MPDATA: A positive definite solver for geophysical flows. – *J. Comput. Phys.* **140**, 459–480.
- SMOLARKIEWICZ, P., R. SHARMAN, J. WEIL, S. PERRY, D. HEIST, G. BOWKER, 2007: Building resolving large-eddy simulations and comparison with wind tunnel experiments. – *J. Comput. Phys.* **227**, 633–653.
- VALLIS, G.K., 2006: *Atmospheric and oceanic fluid dynamics: fundamentals and large-scale circulation*. – Cambridge University Press.
- VETTIN, F., 1857: Über den aufsteigenden Luftstrom, die Entstehung des Hagels und der Wirbel-Stürme. – *Ann. Phys. Chem.* **102**, 246–255.
- VINCZE, M., U. HARLANDER, T. VON LARCHER, C. EGBERS, 2014: An experimental study of regime transitions in a differentially heated baroclinic annulus with flat and sloping bottom topographies. – *Nonlinear Proc. Geoph.* **21**, 237–250.
- VON LARCHER, T., A. DÖRNBRACK, 2014: Numerical simulations of baroclinic driven flows in a thermally driven rotating annulus using the immersed boundary method. – *Meteorol. Z.* **23**, 599–610, DOI: [10.1127/metz/2014/0609](https://doi.org/10.1127/metz/2014/0609)
- VON LARCHER, T., C. EGBERS, 2005: Experiments on transitions of baroclinic waves in a differentially heated rotating annulus. – *Nonlinear Proc. Geoph.* **12**, 1033–1041.
- VON LARCHER, T., A. FOURNIER, R. HOLLERBACH, 2013: The influence of a sloping bottom endwall on the linear stability in the thermally driven baroclinic annulus with a free surface. – *Theo. Comput. Fluid Dyn.* **27**(7), 433–451.
- VON STORCH, H., A. NAVARRA, 1999: *Analysis of climate variability: applications of statistical techniques*. – Springer, 342 pp.
- WILLIAMS, G.P., 1971: Baroclinic annulus waves. – *J. Fluid Mech.* **49**, 417–449.
- WILLIAMS, P.D., P.L. READ, T.W. HAINE, 2010: Testing the limits of quasi-geostrophic theory: application to observed laboratory flows outside the quasi-geostrophic regime. – *J. Fluid Mech.* **649**, 187–203.
- WILLIAMSON, J., 1980: Low-storage Runge-Kutta schemes. – *J. Comput. Phys.* **35**, 48–56.
- YOUNG, R., P. READ, 2008: Breeding and predictability in the baroclinic rotating annulus using a perfect model. – *Nonlinear Proc. Geoph.* **15**, 469–487.
- YOUNG, R., P. READ, 2013: Data assimilation in the laboratory using a rotating annulus experiment. – *Quart. J. Roy. Meteor. Soci.* **139**, 1488–1504.
- ZHU, J., W. RODI, 1992: Computation of axisymmetric confined jets in a diffuser. – *Int. J. Nume. Meth. Fluids* **14**, 241–251.